

Coastal and marine ecosystems: Greenhouse gas sources, sinks and reservoirs: Results from research by the ESSP on coastal and marine ecosystems-related research

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Solich and Ruben Zondervan

UNFCCC-SBSTA meeting Bonn
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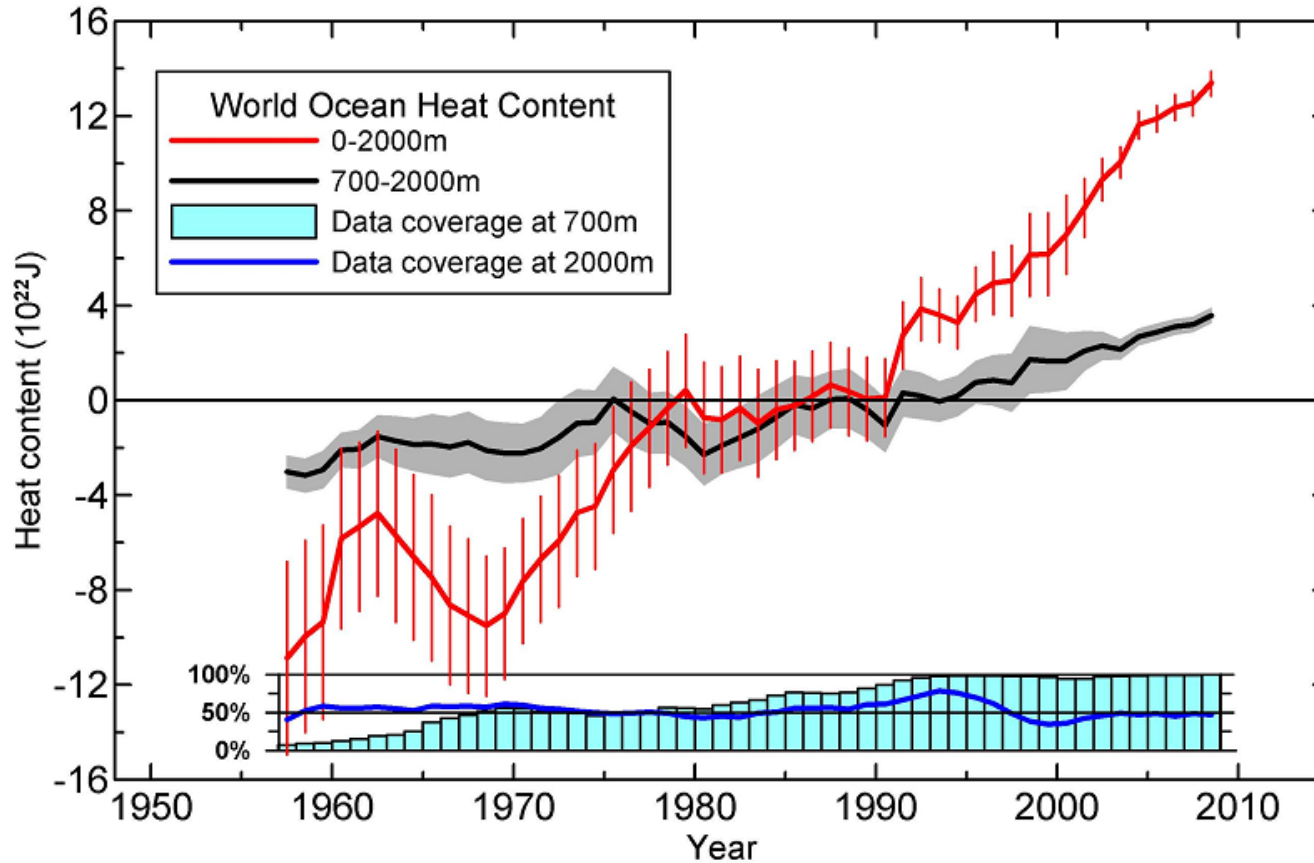


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The oceans have warmed since 1955 due to GHG increase



- A strong positive linear trend is observed in oceans' heat content since 1955
- One third of the warming occurs in the deeper ocean.
- The warming can only be explained by the increase in atmospheric GHG.

Levitus et al. 2012. World ocean heat content and thermosteric sea level change (0-2000), 1955-2010. *GEOPHYSICAL RESEARCH LETTERS*, doi:10.1029/2012GL051106



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Carbon Budget 2010

opinion & comment

CORRESPONDENCE:

Rapid growth in CO₂ emissions after the 2008–2009 global financial crisis

To the Editor — Global carbon dioxide emissions from fossil-fuel combustion and cement production grew 5.9% in 2010, surpassed 9 Pg of carbon (Pg C) for the first time, and more than offset the 1.4% decrease in 2009. The impact of the 2008–2009 global financial crisis (GFC) on emissions has been short-lived owing to strong emissions growth in emerging economies, a return to emissions growth in developed economies, and an increase in the fossil-fuel intensity of the world economy.

Preliminary estimates of global CO₂ emissions from fossil-fuel combustion and cement production show that emissions grew by 0.51 Pg C (5.9%) in 2010 and reached a record high of 9.1±0.5 Pg C (Supplementary Methods). This is the highest total annual growth rate since 2003 (and previously 1979). The 2010 growth overcomes the 1.4% drop in emissions recorded in 2009, which was due to the GFC, putting global CO₂ emissions back on the high-growth trajectory that persisted before the GFC (Fig. 1). Thus, after only one year, the GFC has had little impact on the strong growth trend of global CO₂ emissions that characterized most of the 2000s.

For the past two years (2009 and 2010), emissions growth has been dominated by the emerging economies (Supplementary Table S1). The CO₂ emissions in developed countries (which we take as the Annex B countries from the Kyoto Protocol) decreased 1.3% in 2008 and 7.6% in 2009, but increased 3.4% in 2010, and are now lower than the average emissions during 2000–2007 (Fig. 2). The CO₂ emissions in developing countries (non-Annex B countries) increased 4.6% in 2008, 3.9% in 2009 and 7.6% in 2010; the GFC only causing a 40% decrease in emission growth in 2009 compared with the trend since 2000 (Fig. 2). The 2010 growth was due to high growth rates in a few key emerging economies (Supplementary Table S1) — for example, China 10.4% (0.212 Pg C) and India 9.4% (0.049 Pg C) — although, the contribution from some developed countries was also substantial in absolute terms: for example, United States 4.1% (0.060 Pg C), Russia Federation 5.8%

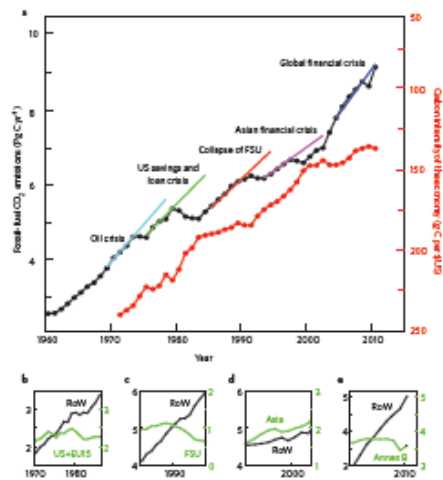


Figure 1 | Global CO₂ emissions and carbon intensity. a, Emissions of CO₂ from fossil-fuel combustion and cement production for the world (Pg C yr⁻¹; black curve) and the carbon intensity of world GDP (g C per \$US 2000); red curve, inverted axis). The most important recent financial crises are highlighted with a linear trend fitted to the five years before the beginning of each crisis. b–e, CO₂ emissions (Pg C) for the regions most affected by each financial crisis (right axis) and the rest of the world (RoW; left axis). b, The oil crisis (1973) and the US savings and loans crisis (1992), where EU15 is the 15 member states of the European Union as of 1993; c, the collapse of the Former Soviet Union (FSU); 1989; d, The Asian financial crisis (1997); e, The recent global financial crisis (2008–2009).

(0.025 Pg C) and the 27 member states of the European Union 2.2% (0.022 Pg C). For recent decades, the growth in global CO₂ emissions can be explained mainly by the growth in economic activity corrected for decreases in the fossil-fuel carbon intensity (FFCI) of the global economy (fossil-fuel and industrial CO₂ emitted per

US dollar of economic output, that is CO₂ per unit of gross domestic product (GDP)¹. Using constant-price GDP measured in purchasing power parities² (real GDP), the FFCI decreased by 1.4% yr⁻¹ on average between 1980 and 2000. Since 2000 however, the FFCI has decreased by only 0.5% yr⁻¹ (Fig. 1), a sign that the positive trend of

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The screenshot shows the website interface for the Carbon Budget 2010. At the top, there are logos for ESSP, DIVERSITAS, IGBP, IHDP, and WCRP. The main header includes the Global Carbon Project logo and a navigation bar with links for Home, Search, Contact Us, Site Map, Carbon Budget, RECCAP, and Urbanization. A sidebar on the left lists categories like Carbon Neutral, About GCP, Activities, Meetings, Publications, Science, Research Programs, and Internet Resources. The main content area features the title 'Carbon Budget' and a large graphic for 'Carbon Budget 2010' with the subtitle 'An annual update of the global carbon budget and trends'. It also indicates the release date as November 2011. Below this, there are sections for 'HIGHLIGHTS' and a grid of links for 'Contributions', 'Presentation', 'Data', 'Policy Brief', 'References', and 'Other Analyses'. A 'Media Information' sidebar on the right provides links for 'Brief Highlights', 'Press Releases', and 'Podcast'.

<http://www.globalcarbonproject.org/carbonbudget>

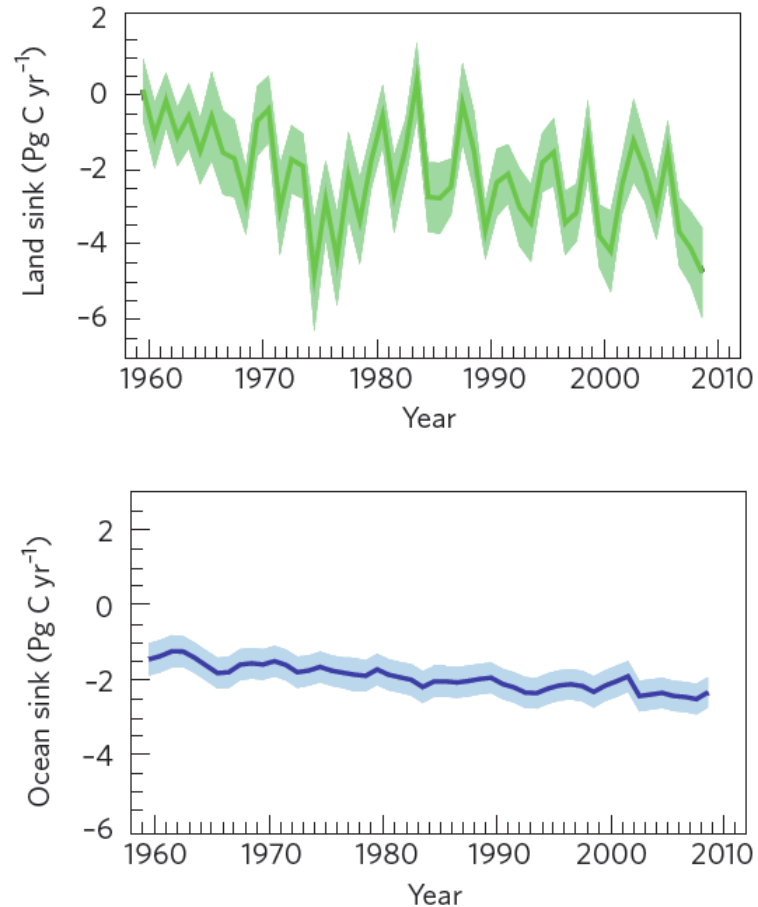
Peters GP, Marland G, Le Quéré C, Boden T, Canadell JG, Raupach MR (2011) Rapid growth in CO₂ emissions after the 2008–2009 global financial crisis. *Nature Climate Change*, doi: 10.1038/nclimate1332.





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Trends in sinks of carbon



nature geoscience FOCUS | PROGRESS ARTICLES
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Trends in the sources and sinks of carbon dioxide

Corinne Le Quééré, Michael R. Raupach, Joseph G. Canadell, Gregg Marland et al.*

Efforts to control climate change require the stabilization of atmospheric CO₂ concentrations. This can only be achieved through a drastic reduction of global CO₂ emissions. Yet fossil fuel emissions increased by 29% between 2000 and 2008, in conjunction with increased contributions from emerging economies, from the production and international trade of goods and services, and from the use of coal as a fuel source. In contrast, emissions from land-use changes were nearly constant. Between 1959 and 2008, 43% of each year's CO₂ emissions remained in the atmosphere on average; the rest was absorbed by carbon sinks on land and in the oceans. In the past 50 years, the fraction of CO₂ emissions that remains in the atmosphere each year has likely increased, from about 40% to 45%, and models suggest that this trend was caused by a decrease in the uptake of CO₂ by the carbon sinks in response to climate change and variability. Changes in the CO₂ sinks are highly uncertain, but they could have a significant influence on future atmospheric CO₂ levels. It is therefore crucial to reduce the uncertainties.

Atmospheric measurements of CO₂ concentration are highly precise and provide an accurate, reliable measure of the increase of CO₂ in the atmosphere every year*. Yet these measurements cannot at present be used to verify global CO₂ emissions estimated from energy data, because the uptake of CO₂ by the land and ocean CO₂ sinks are not quantified with high enough accuracy. Understanding the difference in amount between anthropogenic CO₂ emissions and changes in atmospheric CO₂ concentration requires good estimates of the sinks and good attribution of the causes of changes, both for the emissions and for their partitioning between the natural reservoirs.

Global CO₂ emissions and their partitioning between the atmosphere and the land and ocean CO₂ sinks can be established using a wide range of geophysical and economic data. We have constructed a global CO₂ budget for each year during 1959–2008 and analysed the underlying drivers of each component. The global increase in atmospheric CO₂ was determined directly from measurements. CO₂ emissions from fossil fuel combustion were estimated on the basis of countries' energy statistics. CO₂ emissions from land-use change (LUC) were estimated using deforestation and other land-use data, fire observations from space, and assumptions on the carbon density of vegetation and soils and the fate of carbon. The time evolution of the land and ocean CO₂ sinks, however, cannot be estimated directly from observations. For these terms, we used state-of-the-art models on which we imposed the observed meteorological conditions of the past few decades. The resulting global CO₂ budget provides insight into the global carbon cycle and the emerging trends.

Fossil fuel CO₂ emissions
CO₂ emissions from fossil fuel combustion, including small contributions from cement production and gas flaring, were 8.7±0.5 Pg C yr⁻¹ in 2008, an increase of 2.0% on 2007, 29% on 2000 and 41% above emissions in 1990 (Supplementary Table 1; see Methods). Emissions increased at a rate of 3.4% yr⁻¹ between 2000 and 2008, compared with 1.0% yr⁻¹ in the 1990s (Fig. 1). Emissions continued to track the average of the most carbon-intensive family of scenarios put forward by the Intergovernmental Panel on Climate Change¹ (IPCC) scenario A1FI in Fig. 1a). Since 1990, the growth in fossil fuel CO₂ emissions has been dominated by countries that do not have emissions limitations in the so-called non-Annex B of the Kyoto Protocol (mostly emerging economies in developing countries), where emissions have more than doubled in that time (Fig. 1b). Among Annex B countries (mostly advanced

economies with emissions limitations), growth in some has been offset by declines in others. This recent growth in CO₂ emissions parallels a shift in the largest fuel emission source from oil to coal. Coal contributed 40% of the fossil fuel CO₂ emissions in 2008, compared with 37% for 1990–2000, whereas the contribution of oil changed from 41% for 1990–2000 to 36% in 2008 (Fig. 1c). This shift in the dominant source of fossil fuel emissions has reversed the prevalence of oil since 1960. The growth in emissions since 2000 was also accompanied by an increase in the world per-capita emissions from 1.1 metric tons of carbon in 2000 (Fig. 1d) to an all-time high of 1.3 metric tons of carbon in 2008.

There is growing evidence that the rapid growth in international trade^{2,3} and a shift of Annex B economic activity towards services⁴ were significant in driving non-Annex B CO₂ emission increases due to fossil fuels. Several recent studies provide indicators of the magnitude and time evolution of the share of non-Annex B emissions growth that was due to production of manufactured products exported and consumed in Annex B countries. In 2001, the equivalent of 0.22 Pg C was emitted in non-Annex B countries to produce internationally traded products consumed in Annex B countries⁵. In China alone, 30% of the growth in emissions between 1990 and 2002 was attributable to the production of exports from China that were consumed in other countries⁶, and the share of the growth increased to 50% between 2002 and 2005 (ref. 7). In 1990, 16% of Chinese emissions were from the production of exports, increasing to 30% in 2005. Over half of the exported products were destined for Annex B countries⁸. Complementary studies in some Annex B countries showed that consumption-based emissions (that is, emissions including imported products from non-Annex B countries, but excluding goods and services) were increasing faster than emissions from domestic production⁹. In the UK, for instance, within-country emissions decreased by 5% between 1992 and 2004, whereas consumption-based emissions increased by 12% (ref. 8). In the USA, within-country emissions increased by 6% between 1997 and 2004, whereas consumption-based emissions increased by 17% (ref. 9). In both cases, a key factor driving the growth in consumption-based emissions was the import of manufactured products from China⁸. Taken together, these studies imply that a considerable share of the growth of emissions from non-Annex B countries was associated with international trade. This explained around one-quarter of the growth in non-Annex B emissions since 2000.

The growth in the world gross domestic product (GDP) was a key driver in the recent increase in CO₂ emissions¹⁰. Consequently,

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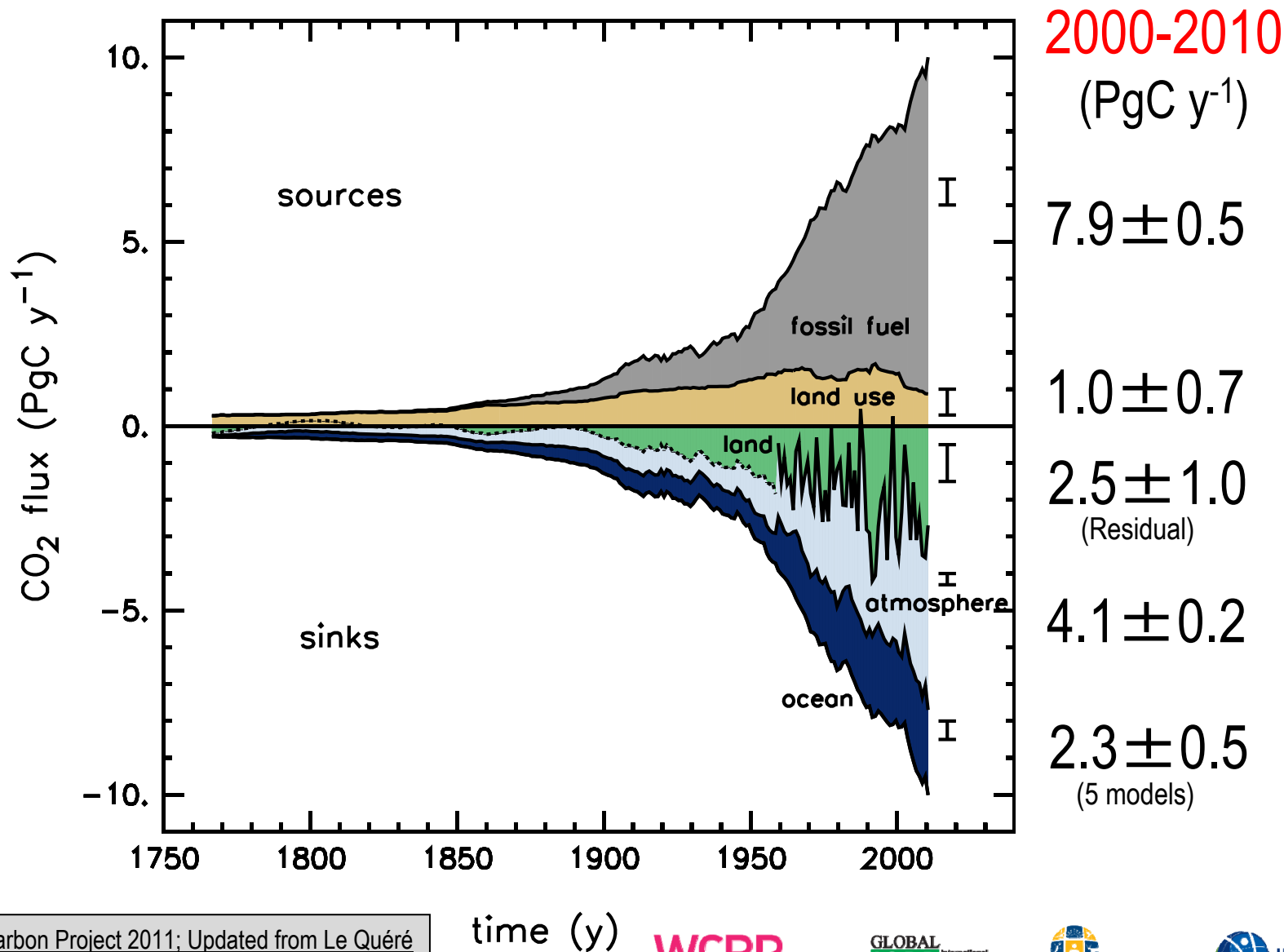
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The fraction of CO₂ emissions that remains in the atmosphere has likely increased from about 40% to 45%. Models suggest that this trend was caused by a decrease in the uptake of CO₂ by carbon sinks in response to climate change and variability.



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Human Perturbation of the Global Carbon Budget



Global Carbon Project 2011; Updated from Le Quéré et al. 2009, Nature G; Canadell et al. 2007, PNAS

time (y)





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Fate of Anthropogenic CO₂ Emissions (2010)

9.1 ± 0.5 PgC y⁻¹



0.9 ± 0.7 PgC y⁻¹ +



5.0 ± 0.2 PgC y⁻¹
50%



2.6 ± 1.0 PgC y⁻¹
26%

Calculated as the residual
of all other flux components



2.4 ± 0.5 PgC y⁻¹
24%

Average of 5 models



Global Carbon Project 2010; Updated from
Le Quéré et al. 2009, Nature Geoscience;
Canadell et al. 2007, PNAS



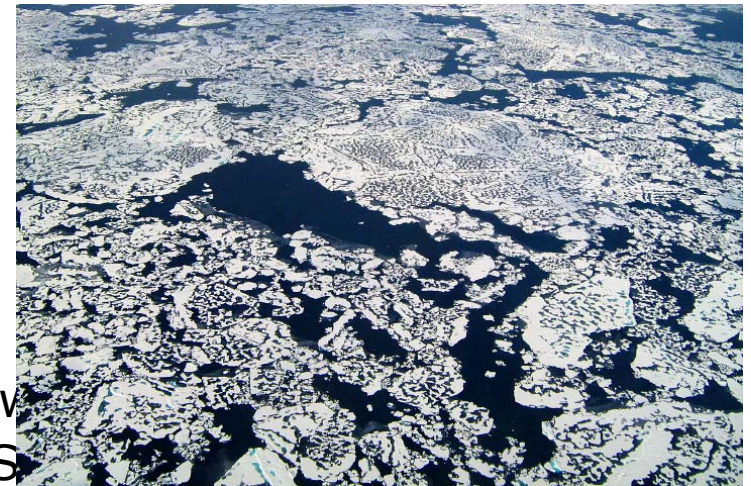


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Large-scale carbon release as warming feedback (generally included in climate-model runs)

- **High northern latitude permafrost thawing** leads to carbon release (CO₂ and methane)
 - **By 2100** under high business-as-usual: release of carbon estimated at 120-420 GtCO₂ → 0.04-0.23° C additional warming
- **“Frozen methane” below sea floor**: possible **destabilization** due to warming and be released (recorded in geological past)
 - Could lead to a slow, chronic release of methane from ocean hydrates
 - blocking warming to return to lower levels for millennia

An observation: Kort et al (2012) measured increased methane levels while flying at low altitudes above the Chukchi and Beaufort S



Schneider von Deimling et al (2011); Schuur and Abbott (2011); Archer et al (2009), Kort, E.A., et al (2012) Atmospheric observations of Arctic Ocean methane emissions up to 82° north. Nature Geoscience 5, 318–321



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ESSP-GWSP on vulnerability of rivers and water security



ARTICLE

Global threats to human water security and river biodiversity

C. J. Vorosmarty¹, P. B. McIntyre^{2,4}, M. O. Gessner³, D. Dudgeon⁴, A. Prusevich⁵, P. Green³, S. Glidden³, S. E. Bunn⁶, C. A. Sullivan⁷, C. Rody Liermann⁸ & P. M. Davies⁹

Protecting the world's freshwater resources requires diagnosing threats over a broad range of scales, from global to local. Here we present the first worldwide synthesis to jointly consider human and biodiversity perspectives on water security using a spatial framework that quantifies multiple stressors and accounts for downstream impacts. We find that nearly 80% of the world's population is exposed to high levels of threat to water security. Massive investment in water technology enables rich nations to offset high stressor levels without remedying their underlying causes, whereas less wealthy nations remain vulnerable. A similar lack of precautionary investment jeopardizes biodiversity, with habitats associated with 65% of continental discharge classified as moderately to highly threatened. The cumulative threat framework offers a tool for prioritizing policy and management responses to this crisis, and underscores the necessity of limiting threats at their source instead of through costly remediation of symptoms in order to assure global water security for both humans and freshwater biodiversity.

Water is widely regarded as the most essential of natural resources, yet freshwater systems are directly threatened by human activities^{1,2} and stand to be further affected by anthropogenic climate change^{3,4}. Water systems are transformed through widespread land cover change, urbanization, industrialization and engineering schemes like reservoirs, irrigation and interbasin transfers that maximize human access to water^{5,6}. The benefits of water provision to economic productivity^{7,8} are often accompanied by impairment to ecosystems and biodiversity, with potentially serious but unquantified costs^{9,10}. Devoting interventions to reverse these trends, including conventions¹¹ and scientific assessments¹² to protect aquatic biodiversity and ensure the sustainability of water delivery systems¹³, requires frameworks to diagnose the primary threats to water security at a range of spatial scales from local to global.

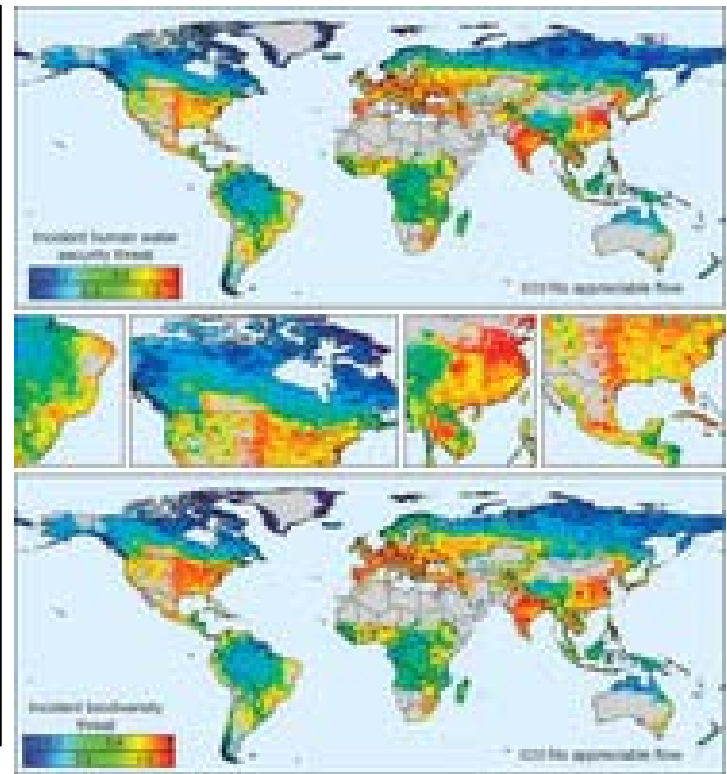
Water issues feature prominently in assessments of economic development¹⁴, ecosystem services¹⁵, and their combination^{16,17}. However, worldwide assessments of water resources¹⁸ rely heavily on fragmented data often expressed as country-level statistics, seriously limiting efforts to prioritize their protection and rehabilitation¹⁹. High-resolution spatial analyses have taken understanding of the human impact on the world's oceans^{20,21} and the human footprint on land²² to a new level, but have yet to be applied to the formal assessment process for freshwater resources²³ despite a recognized need²⁴.

The success of integrated water management strategies depends on striking a balance between human resource use and ecosystem protection^{25,26}. To test the degree to which this objective has been advanced globally, and to assess its potential value in the future, requires systematic accounting. An important first step is to develop a spatial picture of contemporary incident threats to human water security and biodiversity, where the term 'incident' refers to exposure to a diverse array of stressors at a given location. Many stressors threaten human water security and biodiversity through similar pathways, such as pollution, but they also influence water systems in distinct ways. Reservoirs, for example, convey few negative effects on human water supply, but substantially impact on aquatic biodiversity by impeding the movement of organisms, changing flow regimes and altering habitat. Similarly, non-native species threaten biodiversity but are typically inconsequential to human water security.

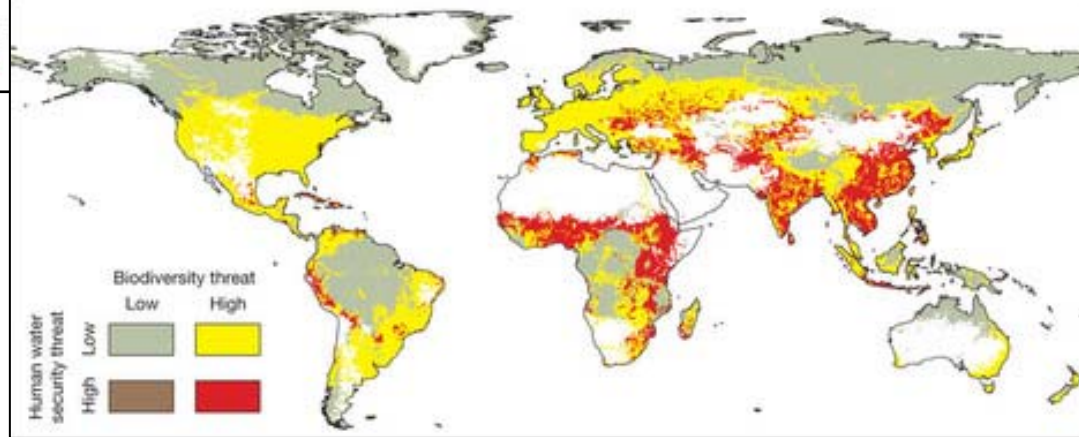
Here we report the results of a global-scale analysis of threats to fresh water that, for the first time, considers human water security and biodiversity perspectives simultaneously within a spatial accounting framework. Our focus is on rivers, which serve as the chief source of renewable water supply for humans and freshwater ecosystems²⁷. We use river networks to redistribute the distinctive impact of stressors on human water security and biodiversity along a continuum from headwaters to ocean, capturing spatial legacy effects ignored by earlier studies. Our framework incorporates all major classes of anthropogenic drivers of stress and enables an assessment of their aggregate impact under often divergent value systems for biodiversity and human water security. Enhancing the spatial resolution by orders-of-magnitude over previous studies (using 30' latitude/longitude grids) allows us to more rigorously test previous assessments on the state of the world's rivers and to identify key sources of threat at sub-national spatial scales that are useful for environmental management. Finally, we make the first spatial assessment of the benefits accrued from technological investments aimed at reducing threats to human water security, revealing previously unrecognized global-scale consequences of local water management practices that are used extensively worldwide.

Global patterns of incident threat

Using a global geospatial framework²⁸, we merged a broad suite of individual stressors to produce two cumulative incident threat indices, one for human water security and one for biodiversity. The resulting



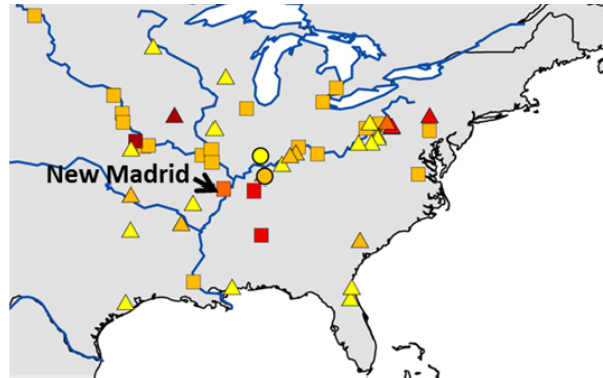
Vorosmarty, C. J., P. B. McIntyre, M. O. Gessner, D. Dudgeon, A. Prusevich, P. Green, S. Glidden, S. E. Bunn, C. A. Sullivan, C. R. Liermann, and P. M. Davies. 2010. Global threats to human water security and river biodiversity. *Nature* 467:555-561.





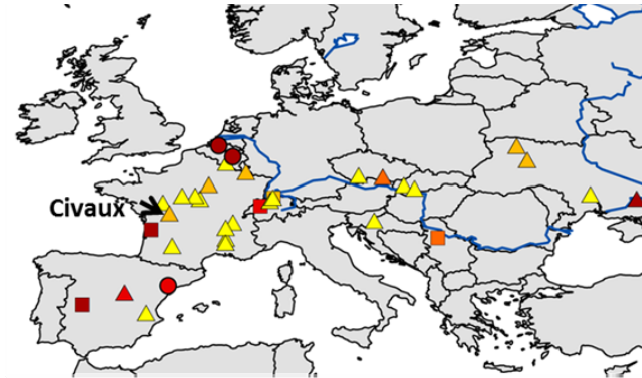
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Vulnerability of US and European electricity supply to climate change



change (%)

- -10 to -5
- < -20
- -20 to -10
- -5 to -1
- -1 to 0



- once through, fresh water
- combination (once-through with supplementary cooling tower)
- △ recirculating with cooling tower(s)



- Thermoelectric (nuclear and fossil fueled) power plants currently produce 91% of all electricity in the US and 78% in Europe
- Directly depend on availability and temperature of water resources for cooling
- Mean capacity decrease of 6.3-19% (Europe) and 4.4-16% (US) depending on cooling system type and climate scenario
- **Strong need for improved climate adaptation strategies in the thermoelectric power sector to assure future energy security.**

van Vliet et al., in press Nature climate change

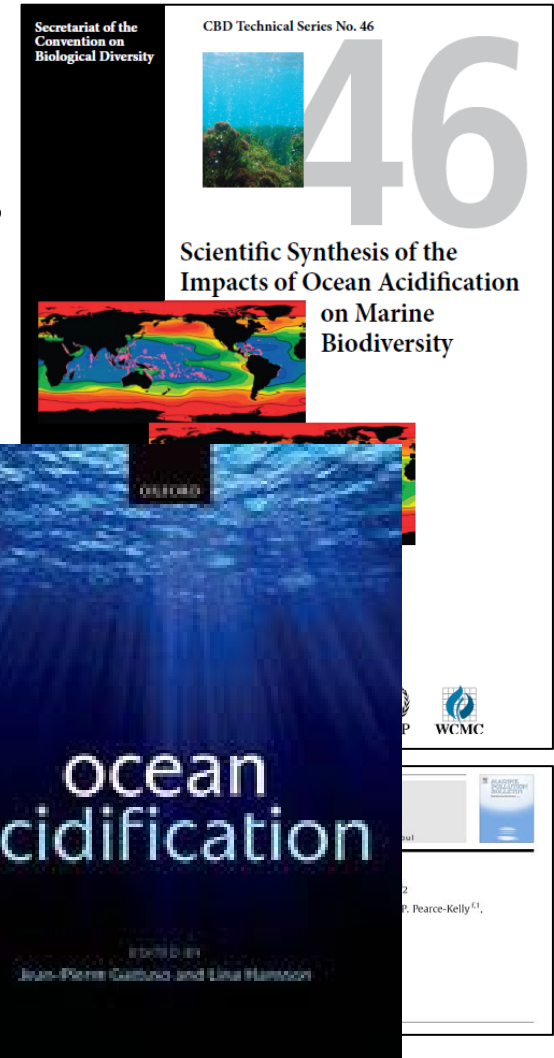
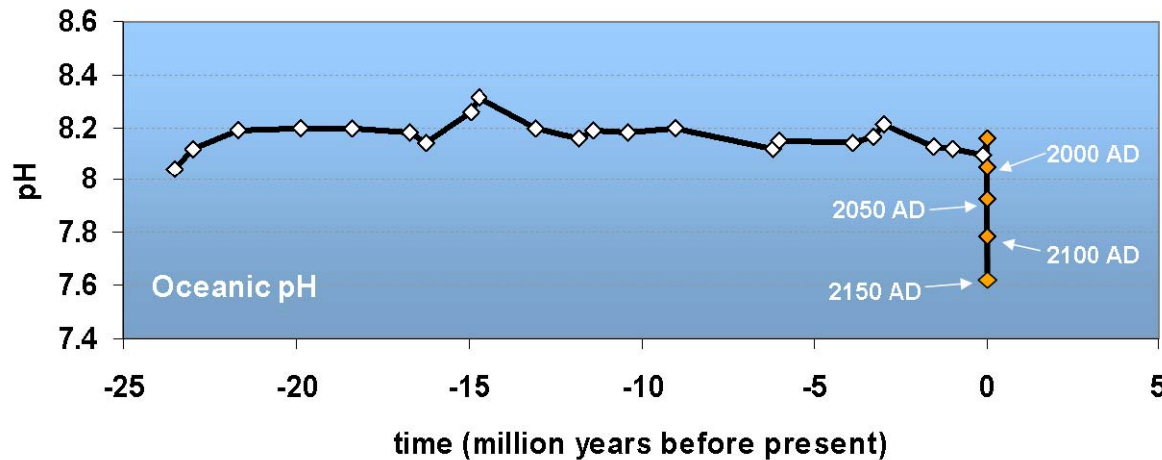


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Coral Reefs: CO₂ and Ocean acidification

Increasing CO₂ concentration acidifying the world oceans

- Likely to have wide ranging adverse effects
- 550 ppm CO₂ coral reefs dissolve (reached by 2050s)
- 450 ppm CO₂ coral stop growing (reached by 2030s)
- Below 350 ppm CO₂ appears to be 'safe'





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Much better insights in ocean acidification impacts

REVIEW

CORRECTED 16 MARCH 2012; SEE LAST PAGE

The Geological Record of Ocean Acidification

Bärbel Hönlisch,^{1,2} Andy Ridgwell,² Daniela N. Schmidt,³ Ellen Thomas,^{4,5} Samantha J. Gibbs,⁶ Apurva Shajr,⁷ Richard Zeebe,⁸ Iac Kump,⁹ Rowan C. Martindale,¹⁰ Sarah E. Greene,^{2,10} Wolfgang Kiessling,¹¹ Justin Ries,¹² James C. Zachos,¹³ Dana L. Royer,¹⁴ Stephen Barker,¹⁴ Thomas M. Marchitto Jr.,¹⁵ Ryan Moyer,¹⁶ Carlos Pelejero,¹⁷ Patricia Ziveri,^{18,19} Gavin L. Foster,²⁰ Bramwen Williams²¹

Ocean acidification may have severe consequences for marine ecosystems; however, assessing its future impact is difficult because laboratory experiments and field observations are limited by their reduced ecologic complexity and sample period, respectively. In contrast, the geological record contains long-term evidence for a variety of global environmental perturbations, including ocean acidification plus their associated biotic response. We review events exhibiting evidence for elevated atmospheric CO₂, global warming, and ocean acidification over the past ~300 million years.

acidification must be unambiguously identified. In recent years, a variety of trace-element and isotopic tools have become available that can be applied to infer past seawater carbonate chemistry. For instance, the boron isotopic composition ($\delta^{11}\text{B}$) of marine carbonate reflects changes in seawater pH, the trace element (such as B, U, and Zn) to-calcium ratio of benthic and planktic foraminifer shells records ambient $[\text{CO}_2]$, and the stable carbon isotopic composition ($\delta^{13}\text{C}$) of organic molecules (alkenones) can be used to estimate surface ocean aqueous $[\text{CO}_2]$.

Because direct ocean geochemical proxy observations are still relatively scarce, past ocean acidification is often inferred from a decrease in the accumulation and preservation of CaCO₃ in marine sediments, potentially indicated by an increased degree of fragmentation of foraminiferal shells (2). However, it is difficult to distinguish

Although similarities exist, no past event perfectly parallels future projections in terms of disrupting the balance of ocean carbonate chemistry—a consequence of the unprecedented rapidity of CO₂ release currently taking place.



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Geochimica et Cosmochimica Acta

Geochimica et Cosmochimica Acta 73 (2009) 2332–2346

www.elsevier.com/locate/gca

Evidence for ocean acidification in the Great Barrier Reef of Australia

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Abstract

Geochemical records preserved in the long-lived carbonate skeleton of corals provide one of the few means to reconstruct

the rapid changes in the reef environment. Here we present a multi-decadal record of $\delta^{11}\text{B}$ and $\delta^{13}\text{C}$ in the past 100 years from the Great Barrier Reef of Australia. The results indicate a significant trend towards acidification in the past decades, which is consistent with a model of ocean acidification driven by increasing atmospheric CO₂. Ocean acidification in the Great Barrier Reef of Australia is likely to have occurred since the late 19th century.

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Increasing calcification rates on reefs off Western Australia contrast with the decline of Australia's Great Barrier Reef and provide additional evidence that recent changes in coral calcification are responses to temperature rather than ocean acidification.

Coral growth is controlled by the physical and chemical properties of the marine environment (1), which are changing rapidly owing to human interference in the global climate system (2–4). Emissions of CO₂ into the atmosphere from the combustion of fossil fuels, deforestation, and altered land use have resulted in current-day atmospheric CO₂ levels of around 390 parts per million (ppm), an increase of about 40% since preindustrial times. Increased concentrations of atmospheric CO₂ are predicted to increase the frequency and severity of mass coral-bleaching events (5). Such changes in the marine environment are, therefore, likely to compromise coral calcification (facilitated by the coral-algal symbi-

Hönisch, B., A. Ridgwell, D. N. Schmidt, E. Thomas, S. J. Gibbs, A. Sluijjs, R. Zeebe, L. Kump, R. C. Martindale, S. E. Greene, W. Kiessling, J. Ries, J. C. Zachos, D. L. Royer, S. Barker, T. M. Marchitto, R. Moyer, C. Pelejero, P. Ziveri, G. L. Foster, B. Williams. 2012. The Geological Record of Ocean Acidification. Science 335:1058-1063.

Wei, G., M. T. McCulloch, G. Mortimer, W. Deng, and L. Xie. 2009. Evidence for ocean acidification in the Great Barrier Reef of Australia. Geochimica et Cosmochimica Acta 73:2332-2346.

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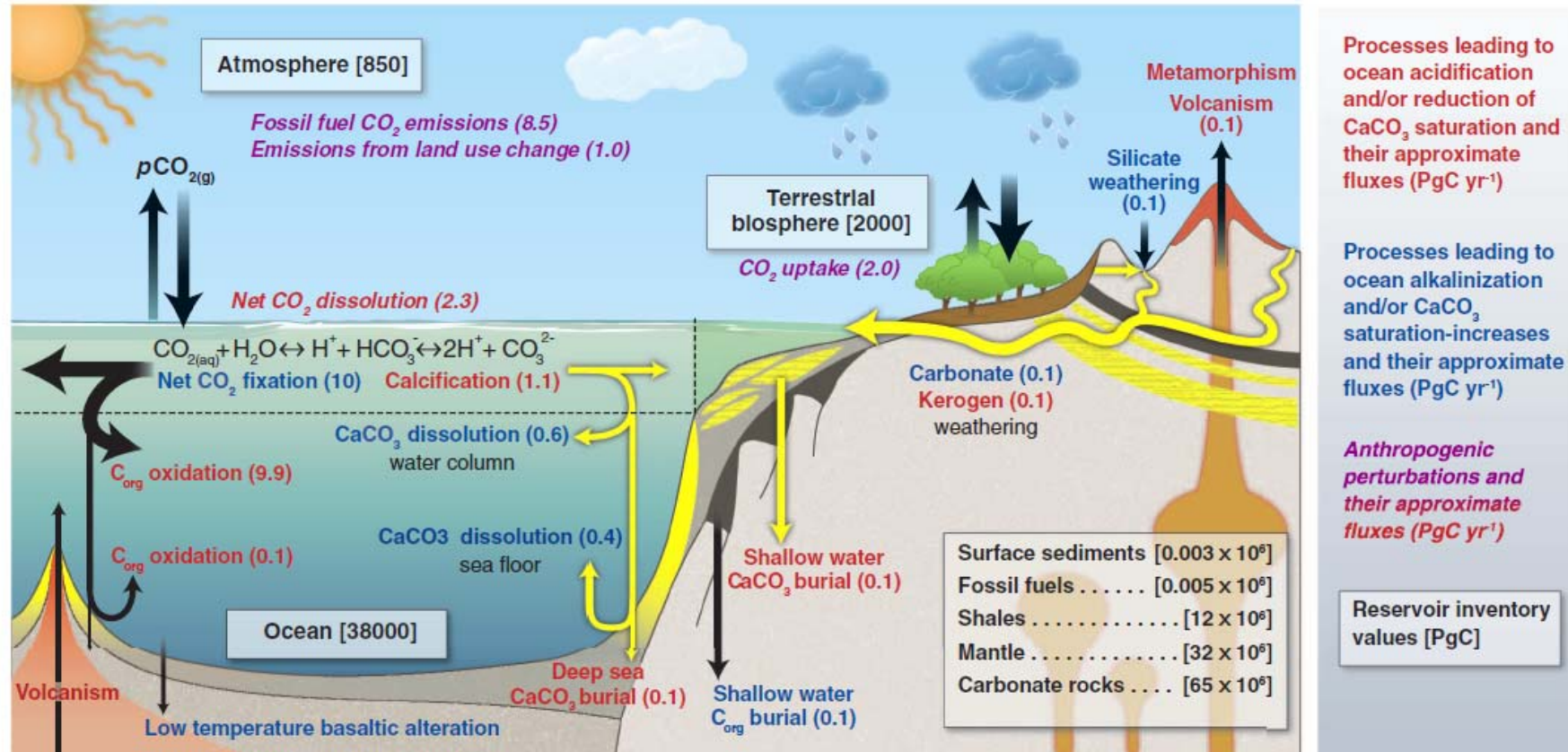


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The process of ocean acidification

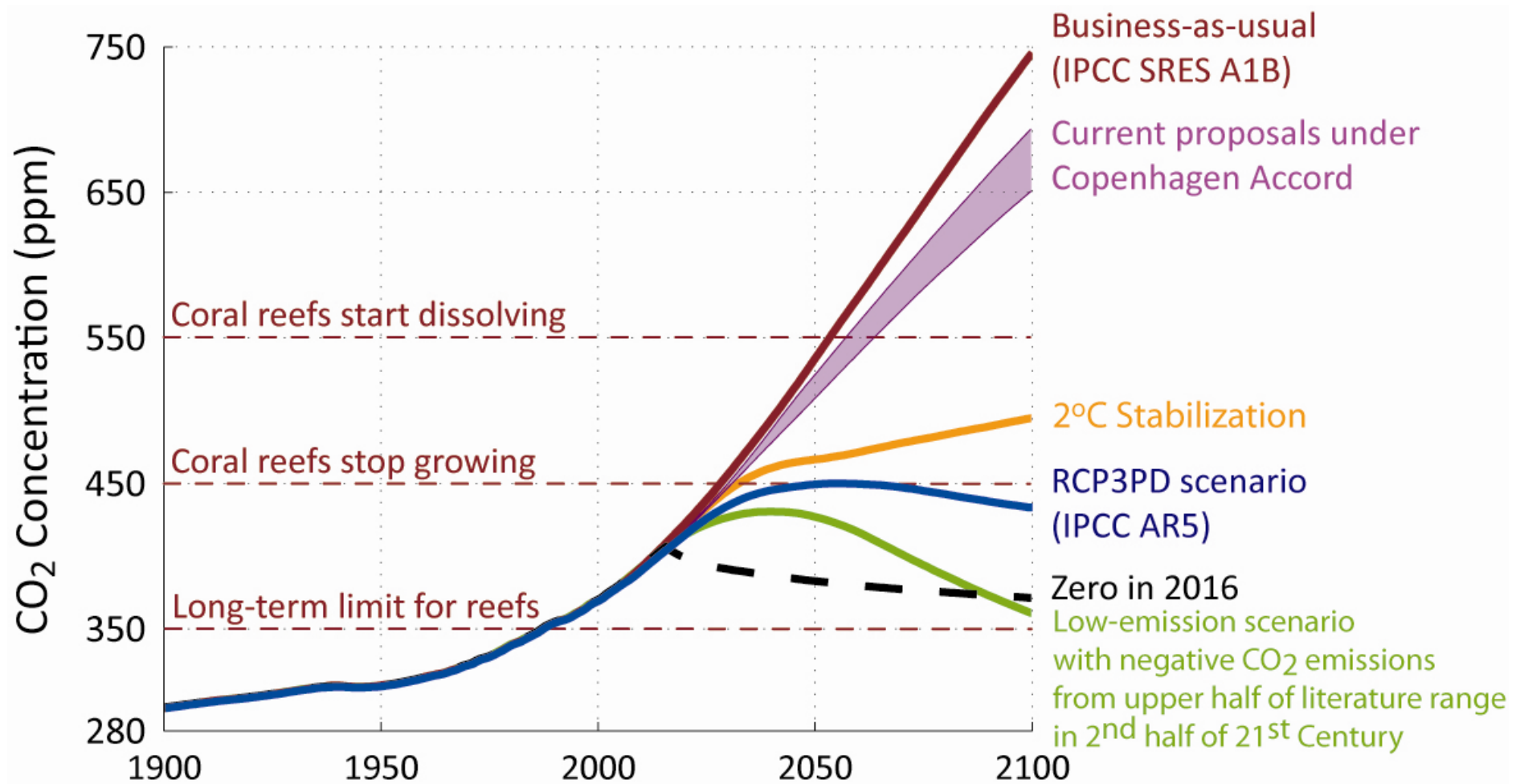


When CO₂ dissolves, it reacts with seawater to form carbonic acid, which then dissociates to bicarbonate, carbonate, and hydrogen ions. The hydrogen ions makes seawater acidic, but this process is buffered on long time scales by the interplay of seawater, seafloor carbonate sediments, and weathering on land. Shown are the major pathways of reduced carbon (**black**) and alkalinity (**yellow**). Ocean acidification or reduction of CaCO₃ saturation are indicated in **red**, and ocean alkalization or CaCO₃ saturation increases are indicated in **blue**.



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Latest scenario projections: Atmospheric CO₂ concentration (ppm)



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