

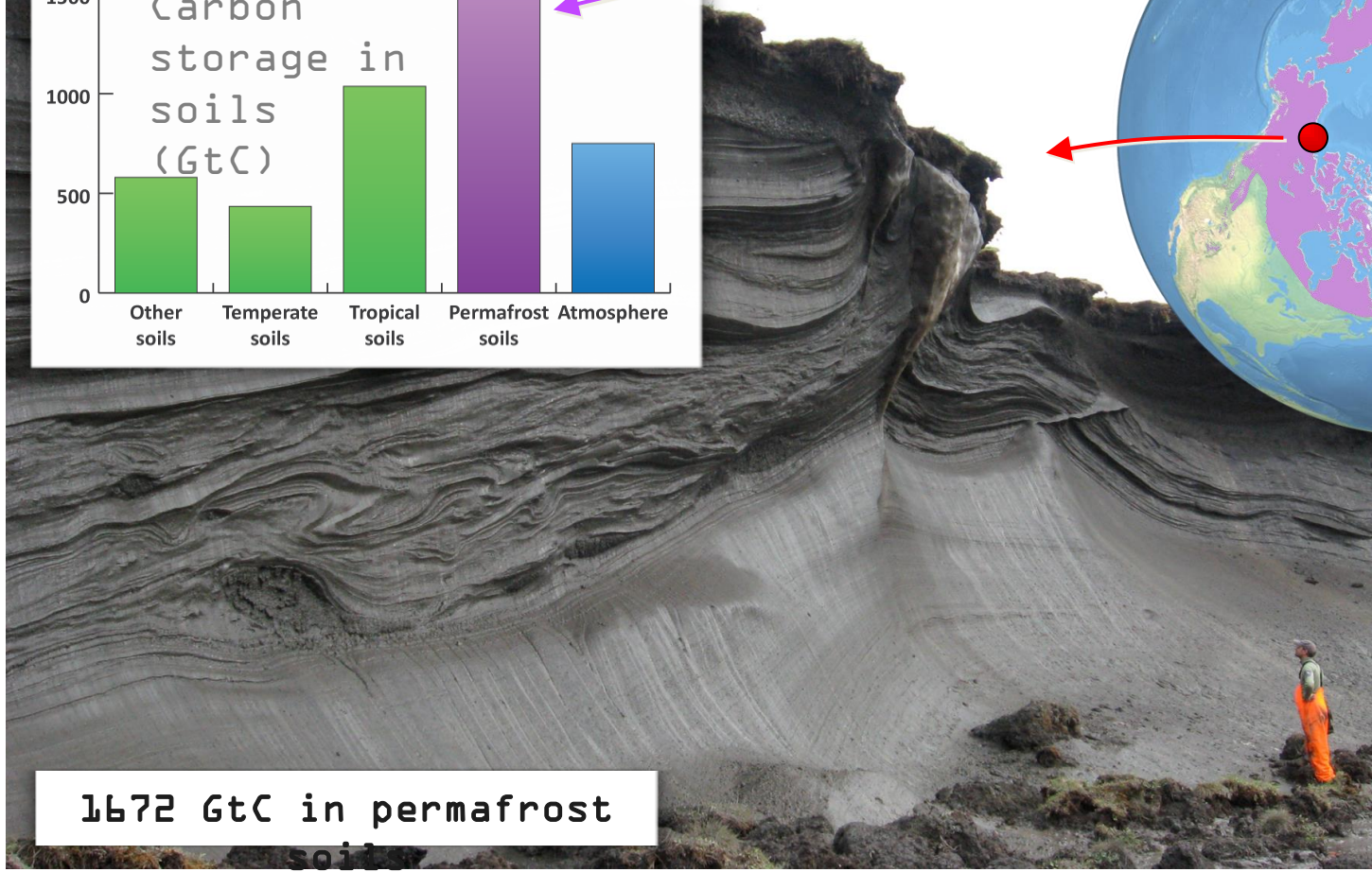
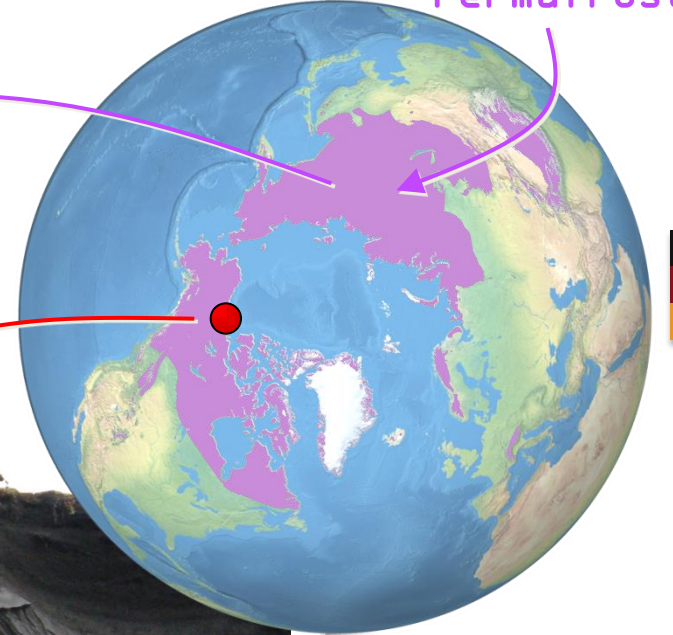
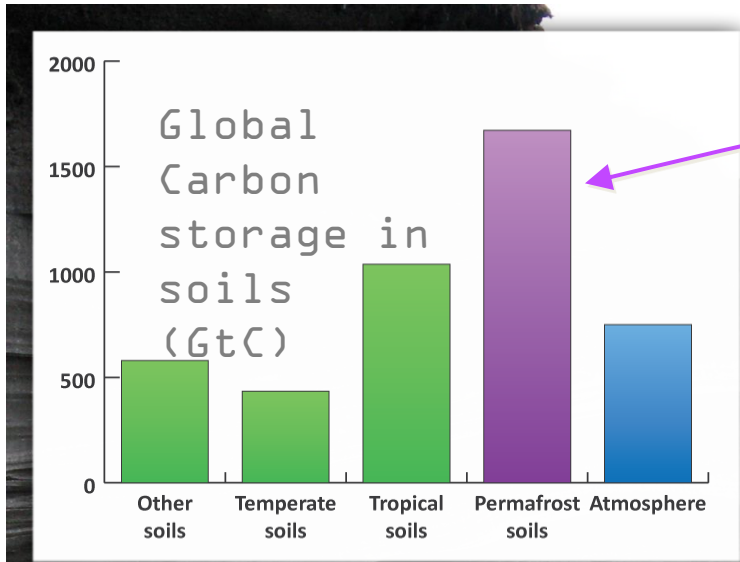
# Permafrost

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with contributions from:  
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Dr. K. Schaefer



# Permafrost



1672 GtC in permafrost soils

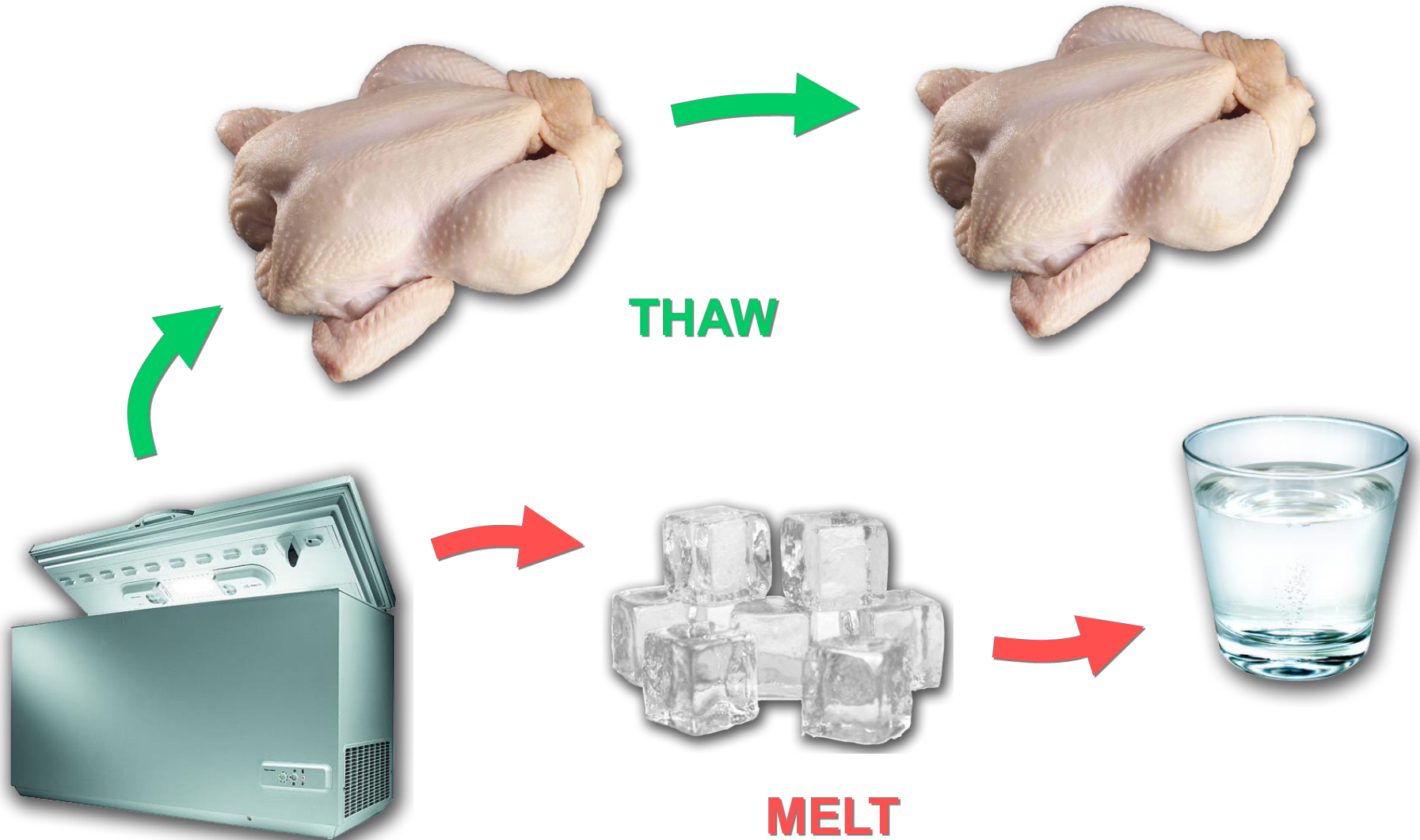
# Permafrost distribution

- Underlies **24%** of the northern hemisphere
- Can be  $\geq$  **2.5 million years** old
- Can be up to **1600 m deep**





# Melt or thaw?



# Permafrost alliance







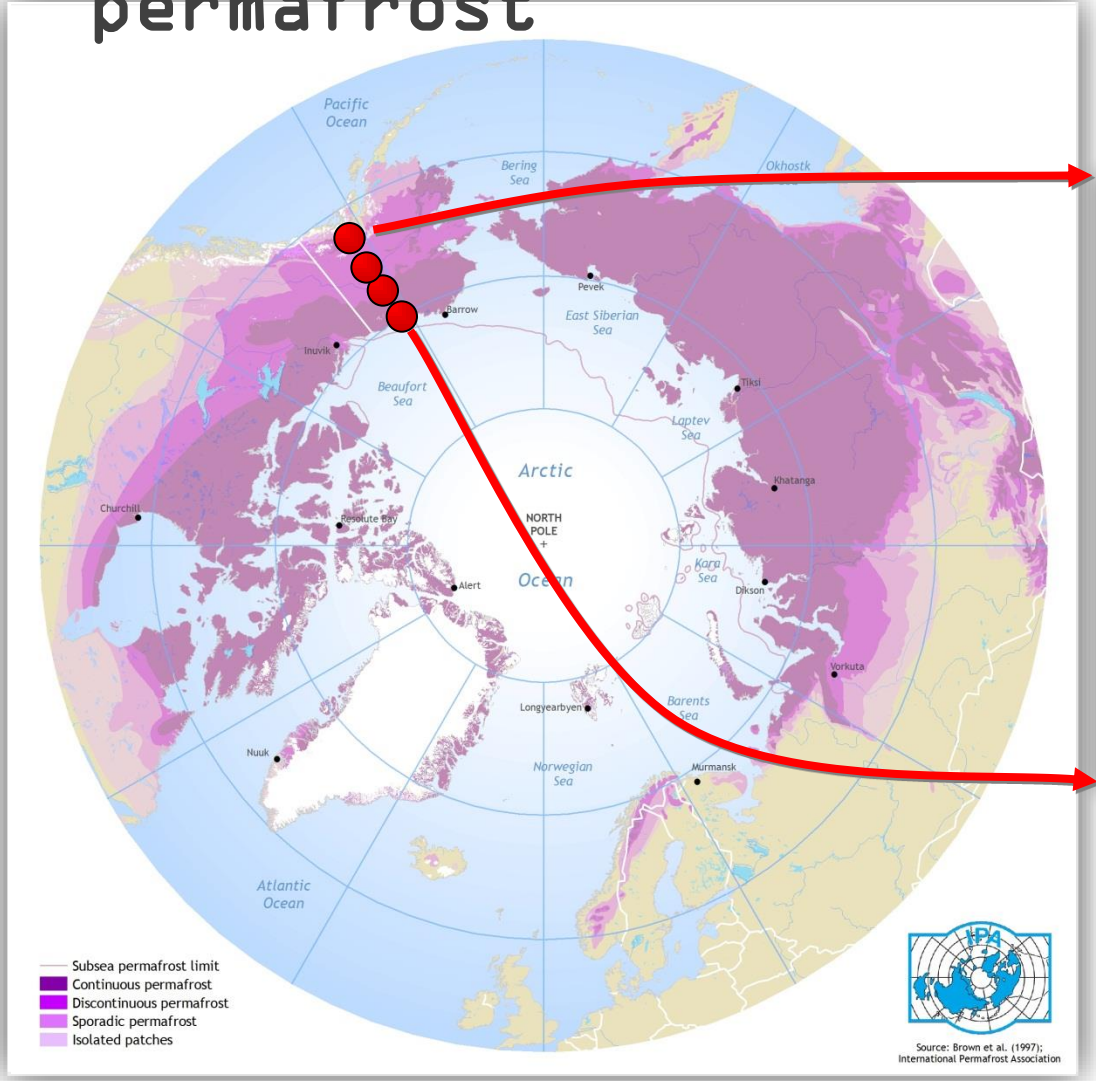


## GRENE-TEA



Source: Brown et al., 1998

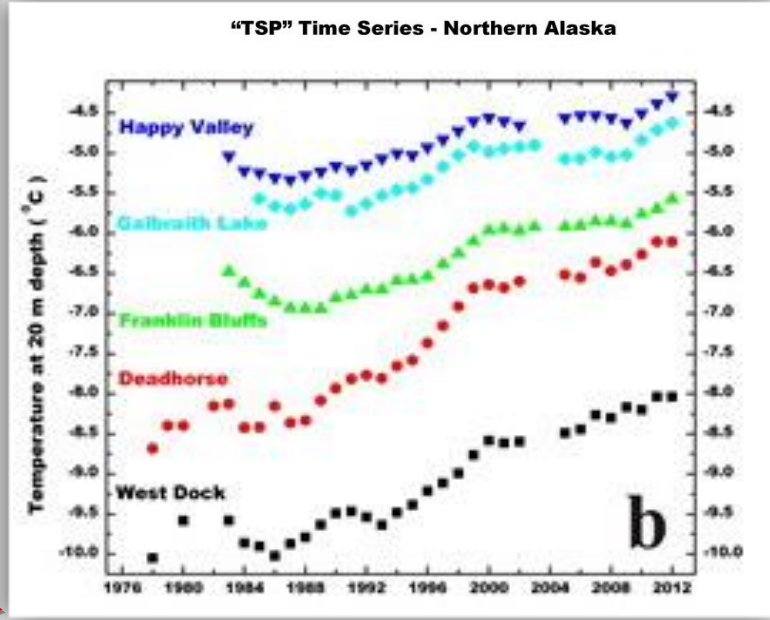
# Observed temperature evolution of Alaska's permafrost



- Subsea permafrost limit
- Continuous permafrost
- Discontinuous permafrost
- Sporadic permafrost
- Isolated patches



Source: Brown et al. (1997); International Permafrost Association

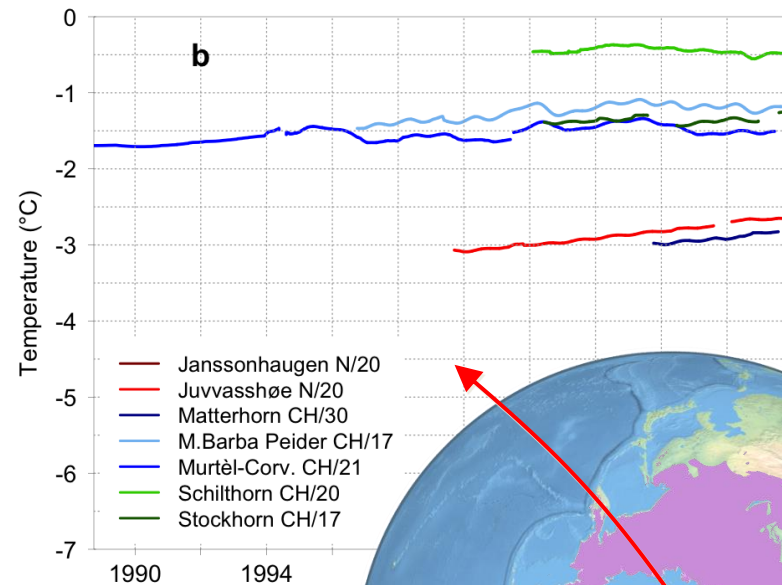
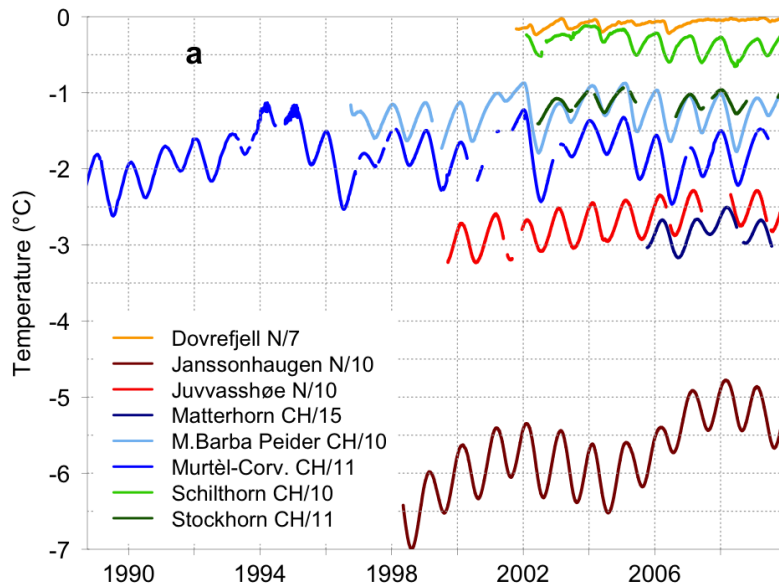


Temperatures at 20 m depth

Romanovsky et al. 2010, Christiansen et al, 2013

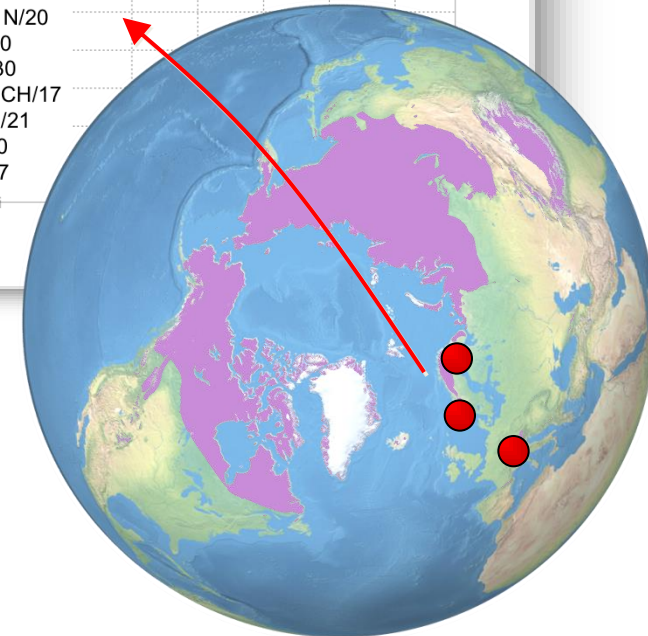


# Mountain permafrost

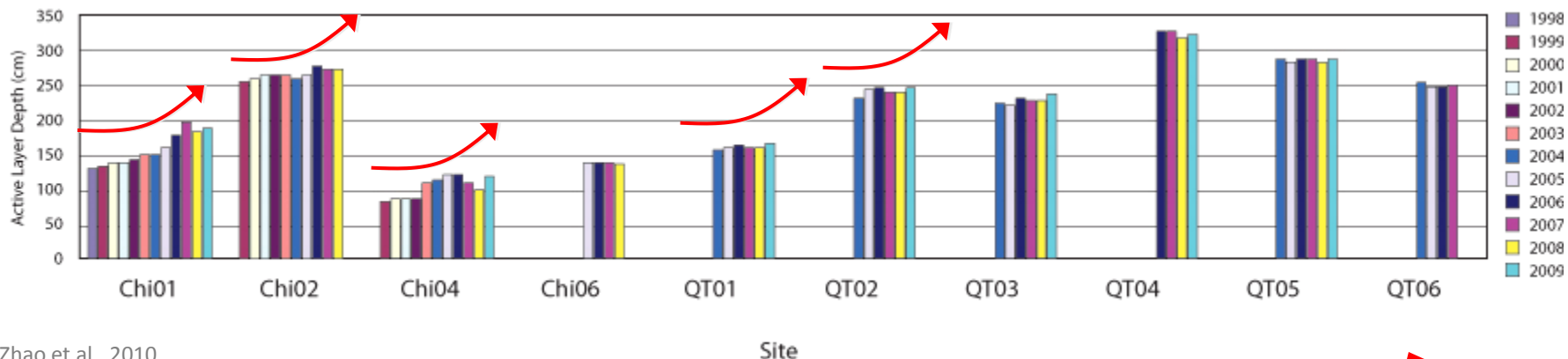


Haeberli et al. 2011

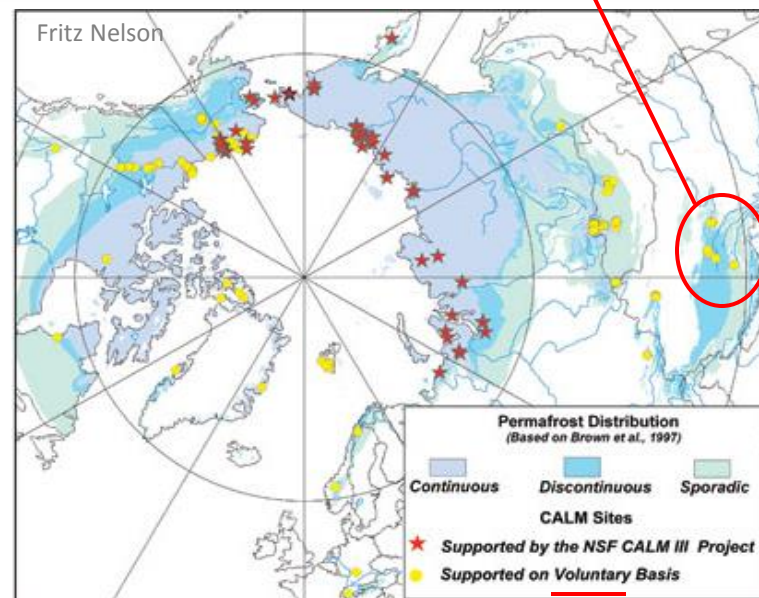
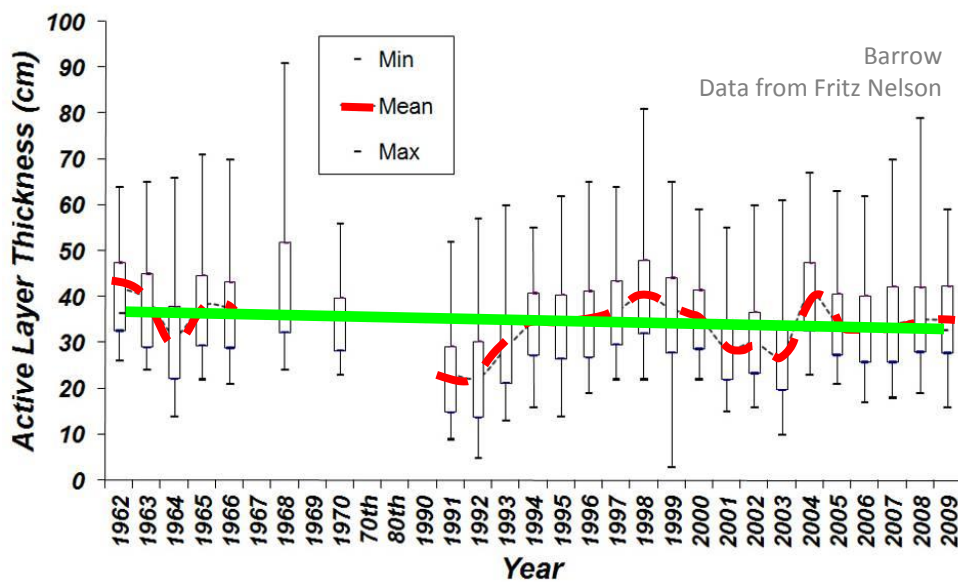
- Mountain permafrost time series are shorter
- Trends difficult to assess because of strong environmental gradients



# Active Layer Depths



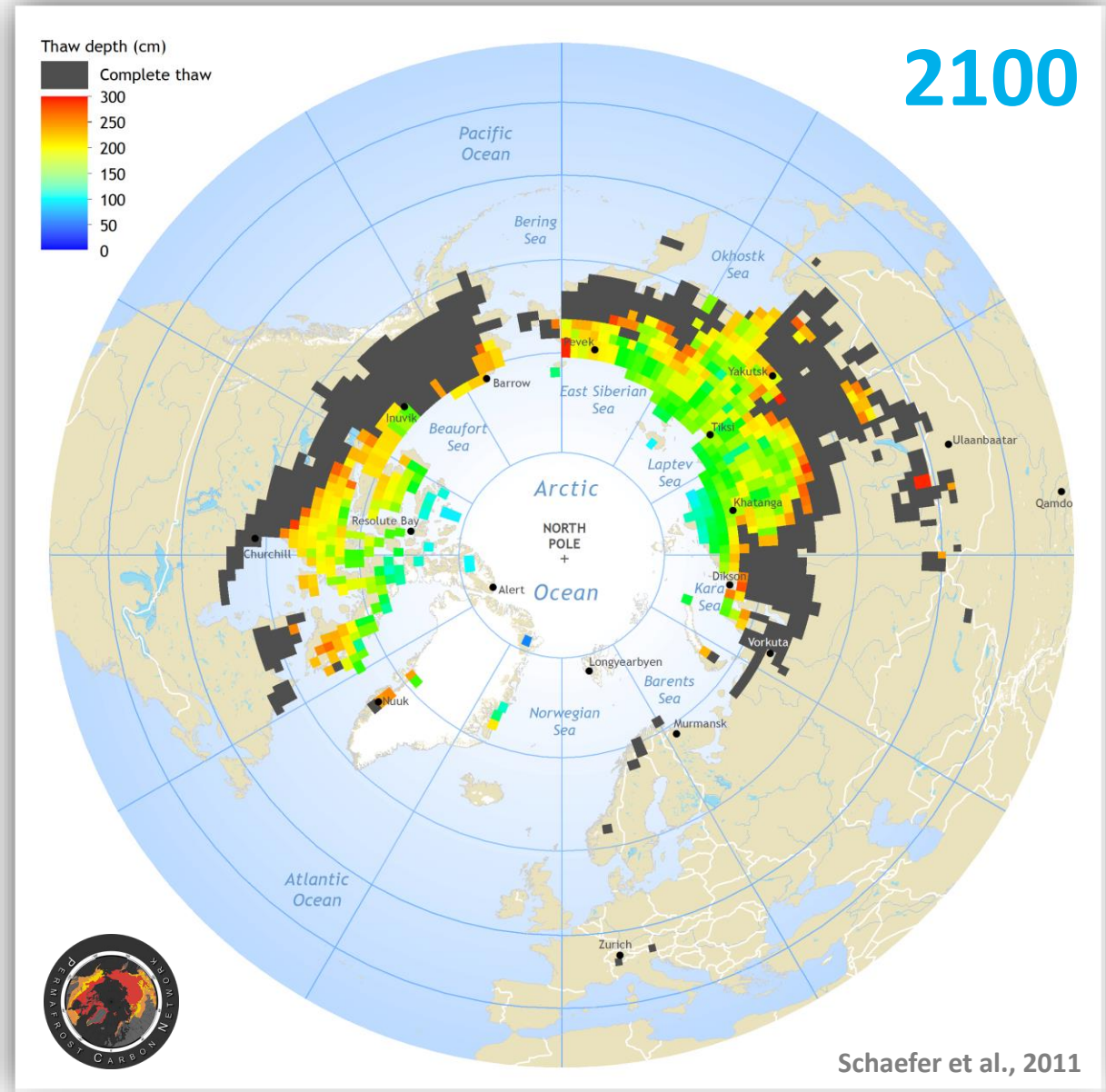
Zhao et al., 2010





# Active Layer

- Thaw in the subarctic
- Deeper active layer in carbon-rich regions

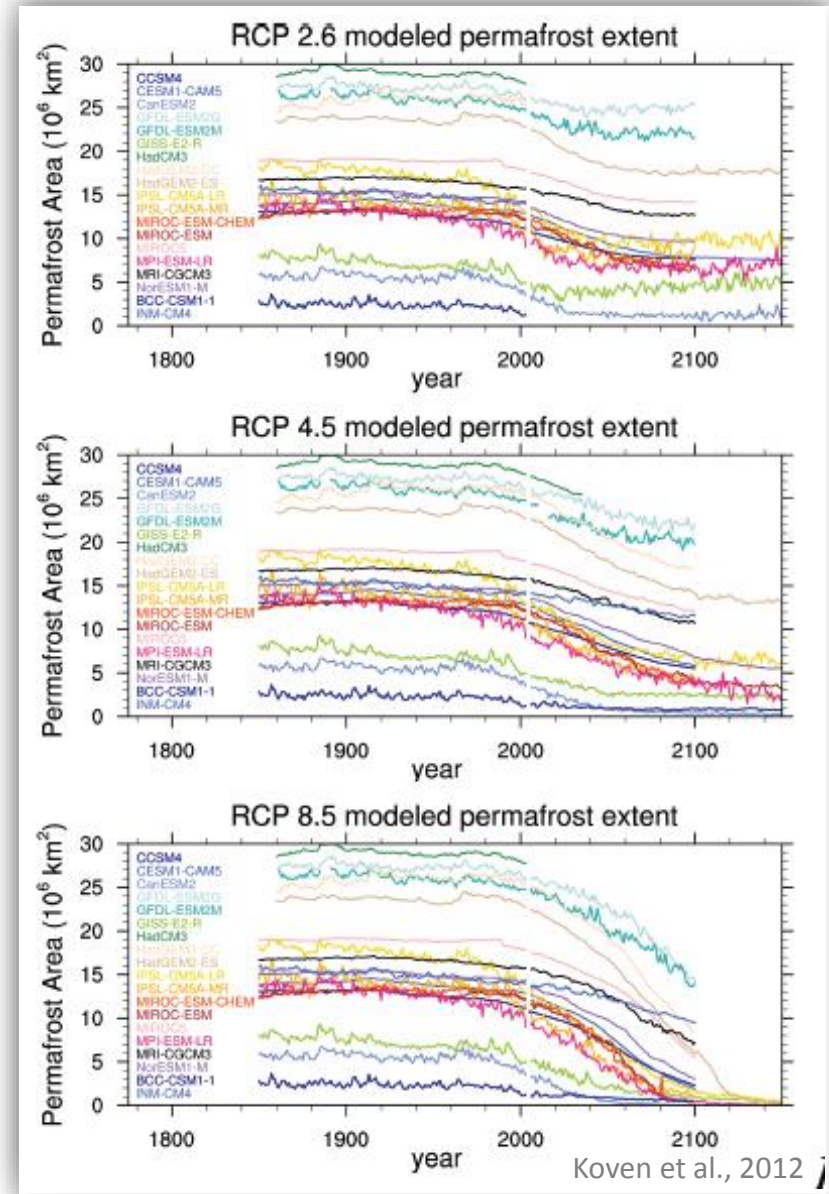


# Thermal modeling



# Permafrost in CMIP5 models

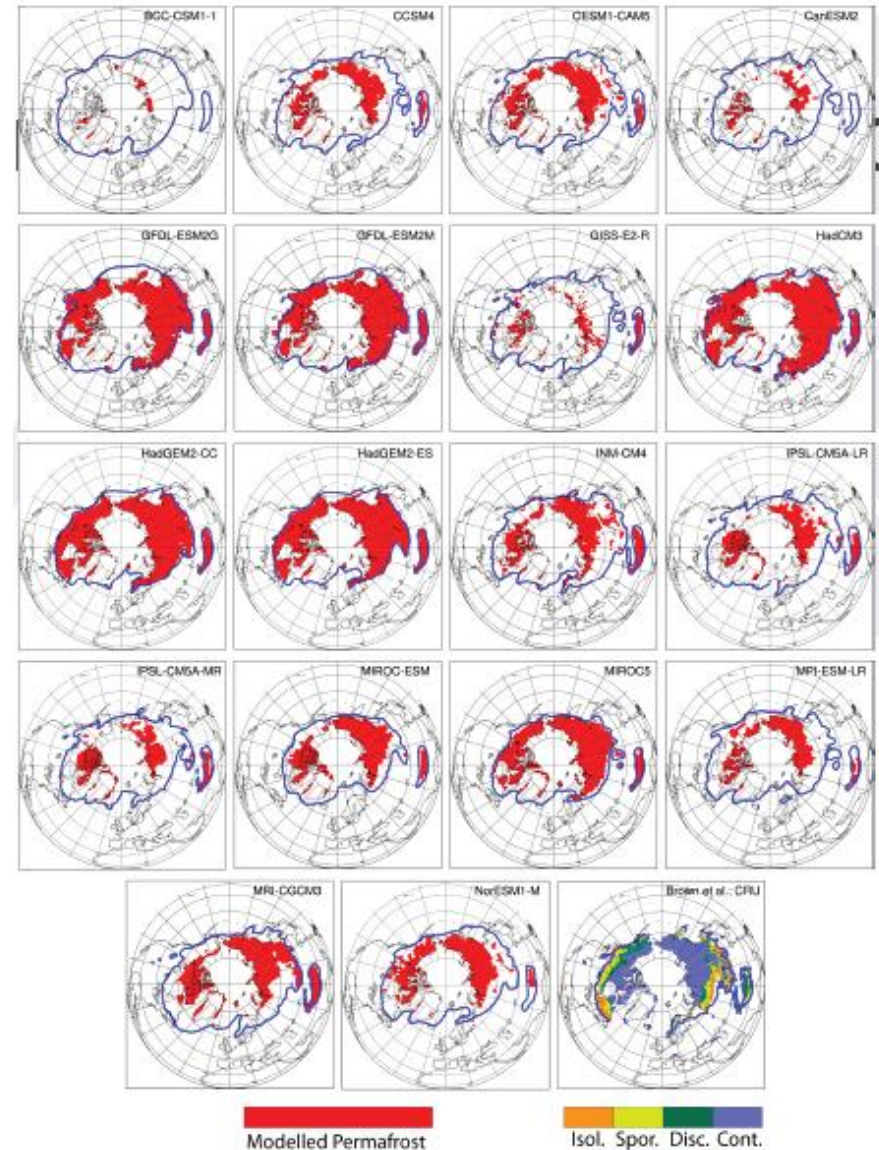
- Large uncertainty on current estimate
- Agreement on decrease in the future (obvious)
- Somehow simulated change is more consistent than simulated present-day distribution...





# Permafrost extent in CMIP5

- The present day estimate remains a major issue
- More model development and evaluation is critical.



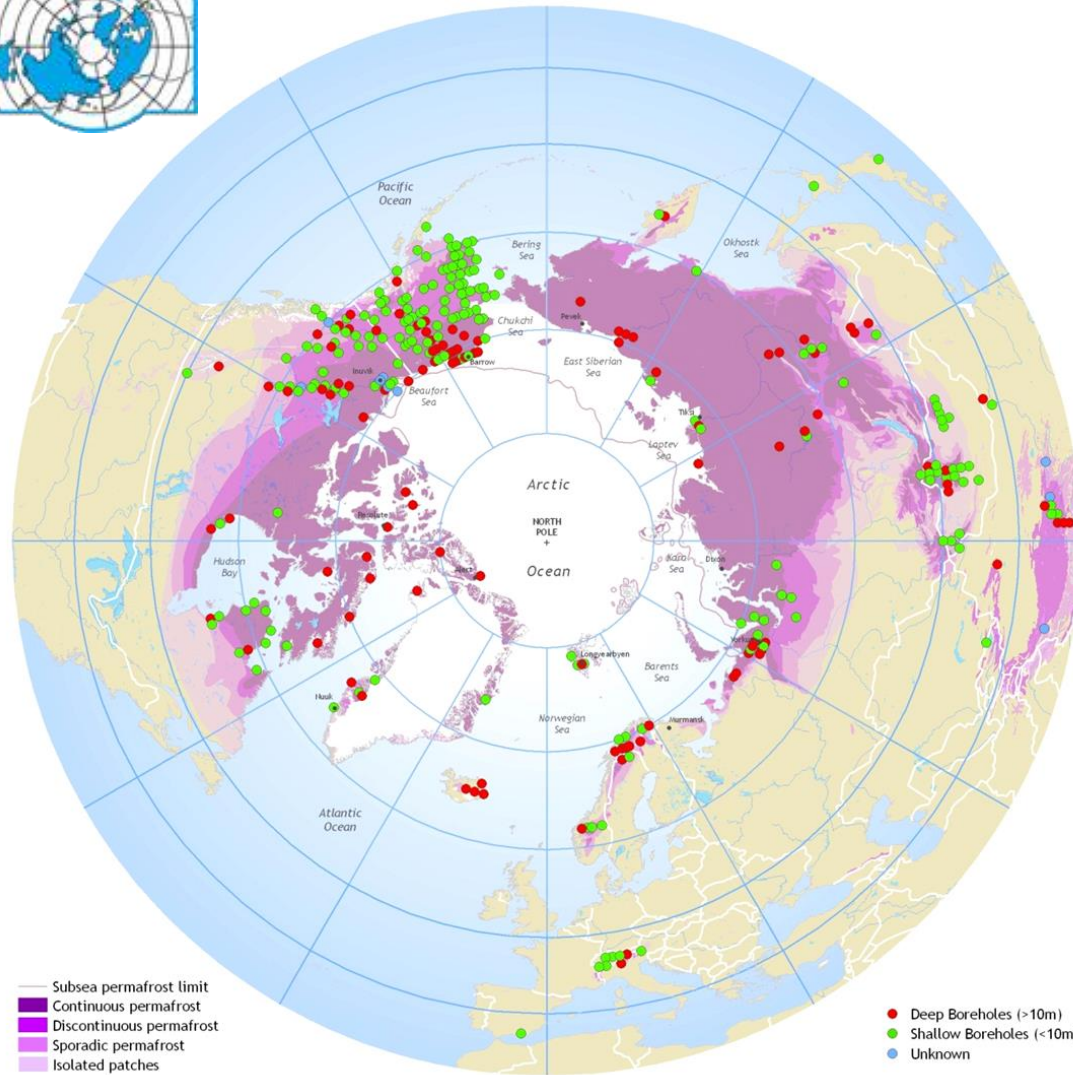
Koven et al., 2012

# Solutions



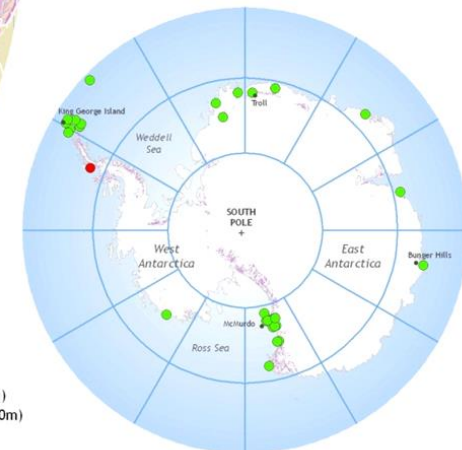


# Borehole distribut



- Subsea permafrost limit
- Continuous permafrost
- Discontinuous permafrost
- Sporadic permafrost
- Isolated patches

- Deep Boreholes (>10m)
- Shallow Boreholes (<10m)
- Unknown



IPA, 2010



# Carbon pools



# Permafrost carbon

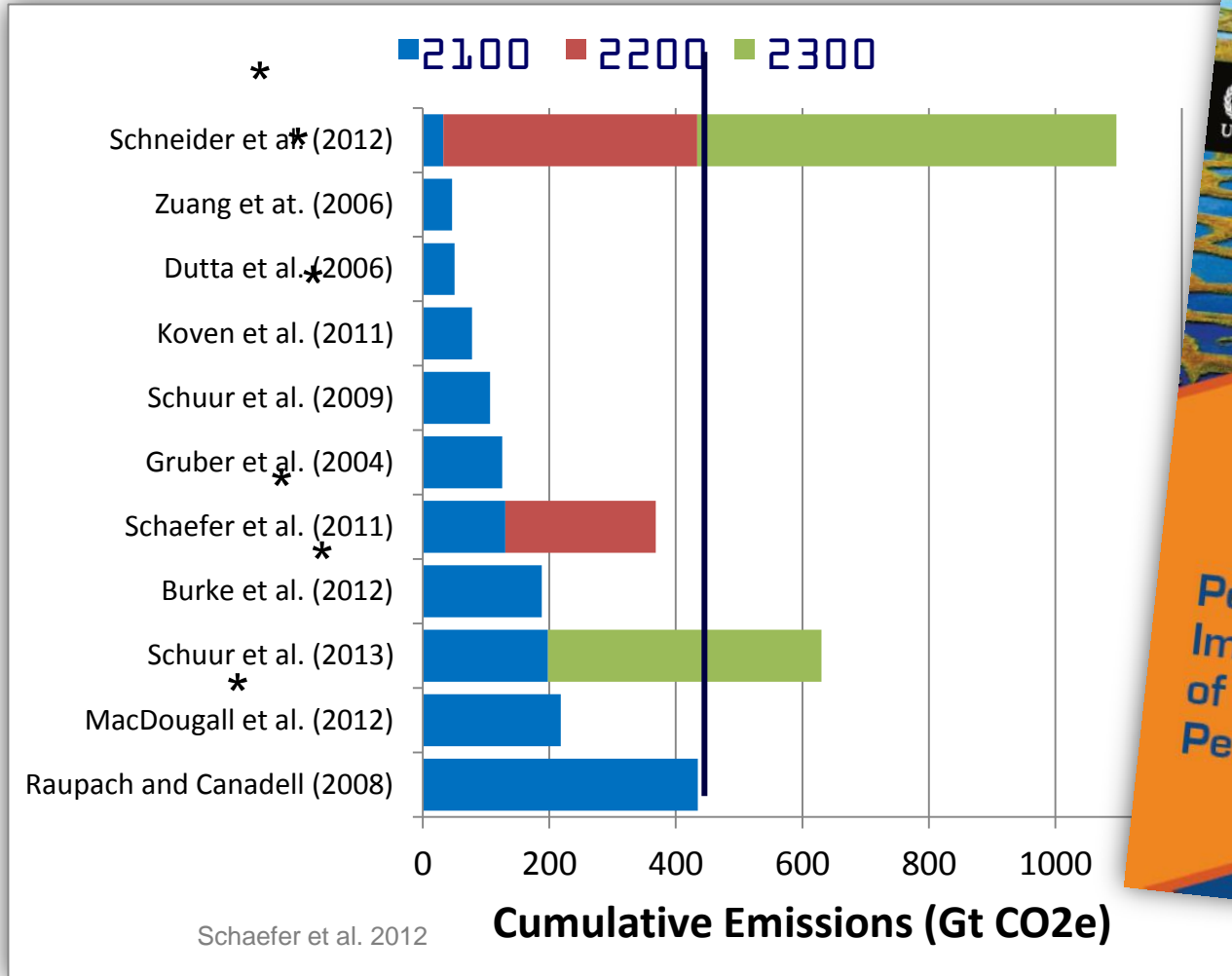
- Fluxes of carbon from thawing permafrost by 2100: between **1.2 and 1.6 Gt C/yr**
- equivalent to **half of all fossil fuel emissions from the dawn of the industrial age to today**

Schaefer et al., 2011

Soil organic carbon content in the upper 100 cm (kg/cm<sup>2</sup>)



# Permafrost carbon emissions



5% to 39% of anthropogenic emissions

\*dynamic model estimates



# Coastal erosion of permafrost



# Coastal erosion of permafrost



## LETTER

### Activation of old carbon by erosion of coastal and subsea permafrost in Arctic Siberia

doi:10.1038/nature11392

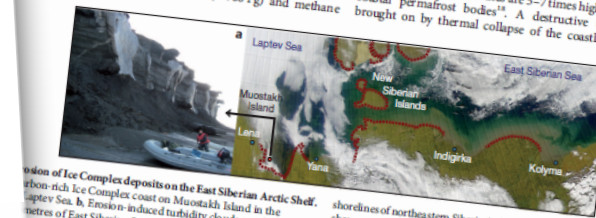
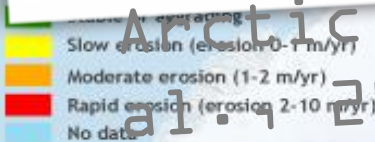
J. E. Vonk<sup>1</sup>\*, L. Sánchez-García<sup>1</sup>\*, B. E. van Dongen<sup>1</sup>, V. Alling<sup>1</sup>, D. Kosmach<sup>2</sup>, A. Charkin<sup>2</sup>, I. P. Semiletov<sup>2,3</sup>, O. V. Dudarev<sup>2</sup>, N. Shakhova<sup>2,3</sup>, P. Roos<sup>4</sup>, T. I. Eglington<sup>5</sup>, A. Andersson<sup>6</sup> & Ö. Gustafsson<sup>7</sup>

The future trajectory of greenhouse gas concentrations depends on interactions between climate and the biogeochemistry<sup>1,2</sup>. Thawing of the Arctic permafrost could release significant amounts of carbon into the atmosphere in this century<sup>3</sup>. Ancient Ice Complex deposits on the East Siberian Arctic Shelf (ESAS)<sup>4,5</sup>, and associated shallow subsea permafrost<sup>6,7</sup>, are two large pools of permafrost carbon, yet their vulnerabilities towards thawing and decomposition are largely unknown<sup>8,9</sup>. Recent Arctic warming is stronger than has been predicted by several degrees, and is particularly pronounced over the coastal ESAS region<sup>10,11</sup>. There is thus a pressing need to improve our understanding of the links between permafrost deposits and climate in this relatively inaccessible region. Here we show that extensive release of carbon from these Ice Complex deposits dominates (57 ± 2 per cent) the sedimentary carbon budget of the ESAS, the world's largest continental shelf, overhauling the marine and topsoil terrestrial carbon pools. Inverse modelling of the dual-carbon isotope composition of organic carbon accumulating in ESAS surface sediments, using Monte Carlo simulations to account for uncertainties, suggests that 4 ± 10 teragrams of old carbon is activated annually from Ice Complex permafrost, an order of magnitude more than has been suggested by previous studies<sup>12</sup>. We estimate that about two-thirds (6 ± 16 per cent) of this old carbon escapes to the atmosphere as methane, with the remainder being re-buried in shelf sediments. Thermal collapse and erosion of these carbon-rich isotone coastline and seafloor deposits may accelerate with the amplification of climate warming<sup>13</sup>. The large magnitude of shallow permafrost carbon pools relative to atmospheric pools of carbon dioxide (~760 Pg) and methane

(~3.5 Pg) suggests that carbon release from thawing permafrost has the potential to affect large-scale carbon cycling. Arctic permafrost can be divided into three main compartments: terrestrial (tundra and taiga permafrost (~1,000 Pg C)), Ice Complex (coastal and inland permafrost (~400 Pg C)) and subsea permafrost (~1,400 Pg C). Even without considering subsea permafrost, the carbon held in the top few metres of the pan-Arctic permafrost complex holds in the half of the global soil organic carbon pool<sup>14</sup>. Investigations of Arctic greenhouse gas releases have focused on terrestrial permafrost systems<sup>15,16</sup>, and only recently on subsea permafrost<sup>17,18,19</sup>, with a notable scarcity of studies on the thawing exposed coastline of the Eastern Siberian Sea (ESS) in particular, the deposits (Fig. 1a). The origin of the ~1-million-km<sup>2</sup> deposits (with average depth 25 m) dominating northeastern Siberia (and parts of Pleistocene material is quite distinct from peat and mineral soil of other Arctic permafrost<sup>20</sup>). These relict soils of the steppe-tundra ecosystem have high carbon contents (1–5%)<sup>21</sup>. The export of organic carbon from the eroding ESAS Ice Complex is presently estimated at 4 Tgyr<sup>-1</sup> (ref. 14), yet it has also been proposed that erosion from the Lena Delta coastline alone might contribute this amount<sup>22</sup>. Clearly, large uncertainties remain regarding the magnitude of erodible carbon export from land to the shelf.

The extensive coastal exposure of the Ice Complex deposits (ICD) makes them potentially more vulnerable than other terrestrial permafrost; ICD retreat rates are 5–7 times higher than those of other coastal permafrost bodies<sup>23</sup>. A destructive thaw-erosion process brought on by thermal collapse of the coastline promotes surface

- 44 Tg of carbon activated by erosion per year (Vonk et al., 2012)
- Between 4 and 46 Tg per year in the Arctic (Lantuit et al., 2013)



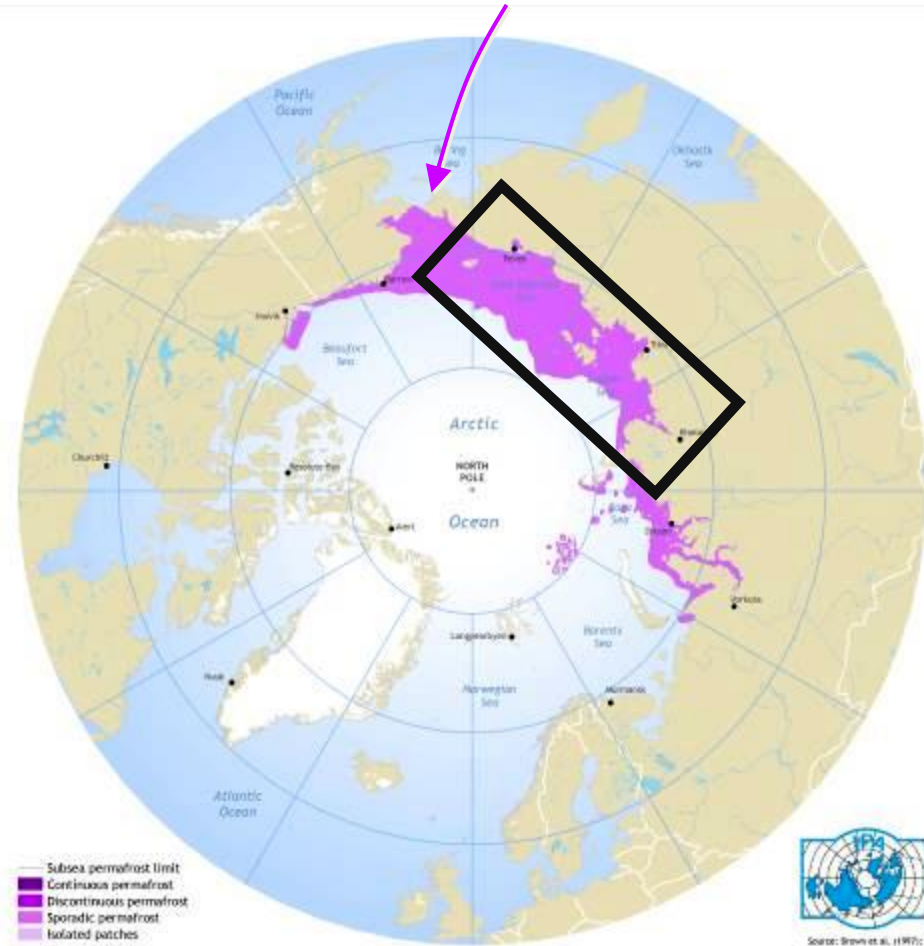
erosion of Ice Complex deposits on the East Siberian Arctic Shelf. a, Carbon-rich Ice Complex coast on Muostakh Island in the Laptev Sea. b, Erosion-induced turbidity clouds envelop several metres of East Siberian Sea coastal waters. Note the rounded shorelines of northeastern Siberia, indicative of coastal erosion. Red dashed line shows areas of intensive ongoing erosion. (Satellite image of 24 August 2000, available at <http://satblear.chm.nyu.edu>.)

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# Gas hydrate



## Subsea permafrost



Source: Brown et al., 1997; International Permafrost Association

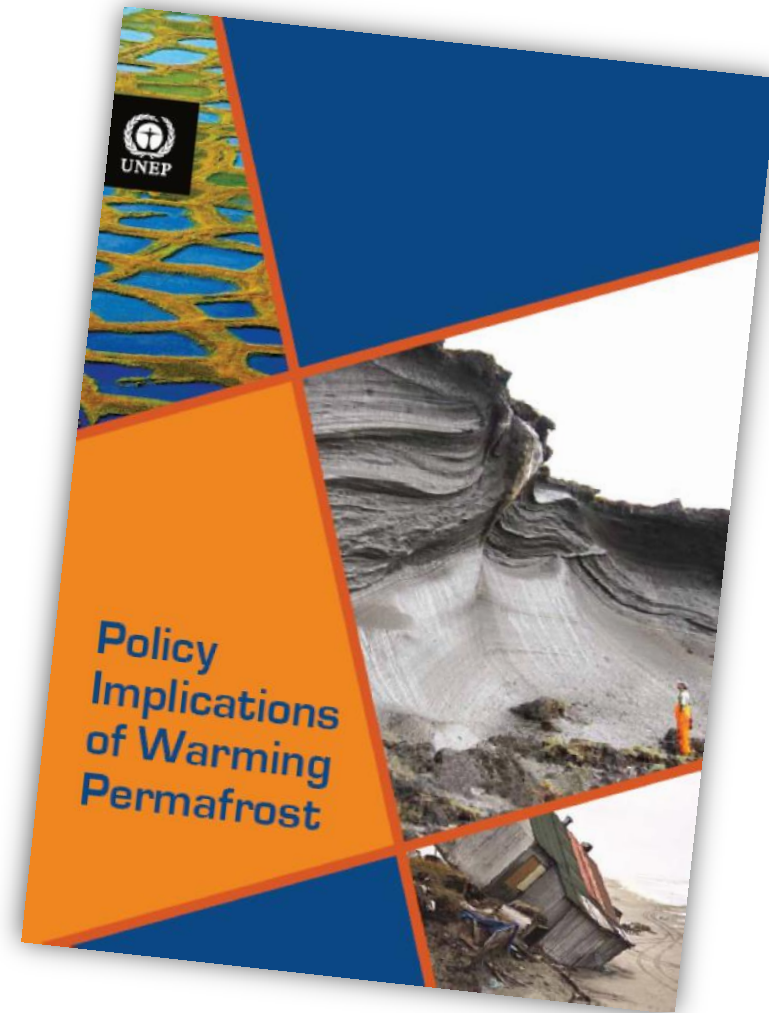
# The way forward





# UNEP recommendations

1. Special IPCC assessment on permafrost emissions
2. National permafrost monitoring networks
3. National Adaptation Plans



# Research needs

1. Model validation
2. Common framework for assessment of carbon bioavailability
3. Lateral fluxes from permafrost areas (coastal and rivers)
4. „between-reservoirs“ degradation of permafrost
5. Subsea permafrost and in situ vs. Gas hydrate methane emission
6. Role of nitrogen and phosphorus

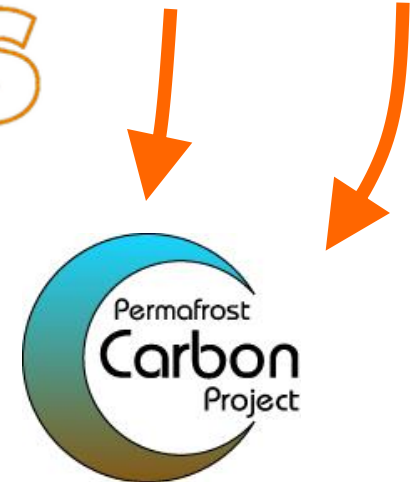
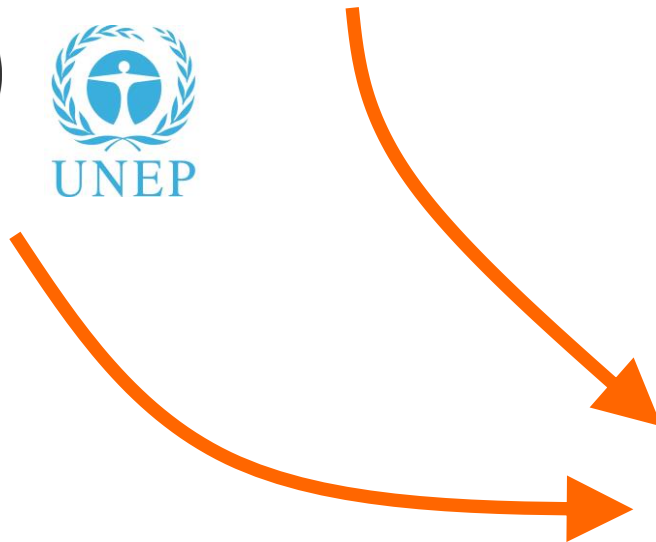
# More collaboration



United Nations  
Framework Convention on  
Climate Change



**WCRP**  
World Climate Research Programme



# www.permafrost.org

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