



# Modelling the global carbon cycle

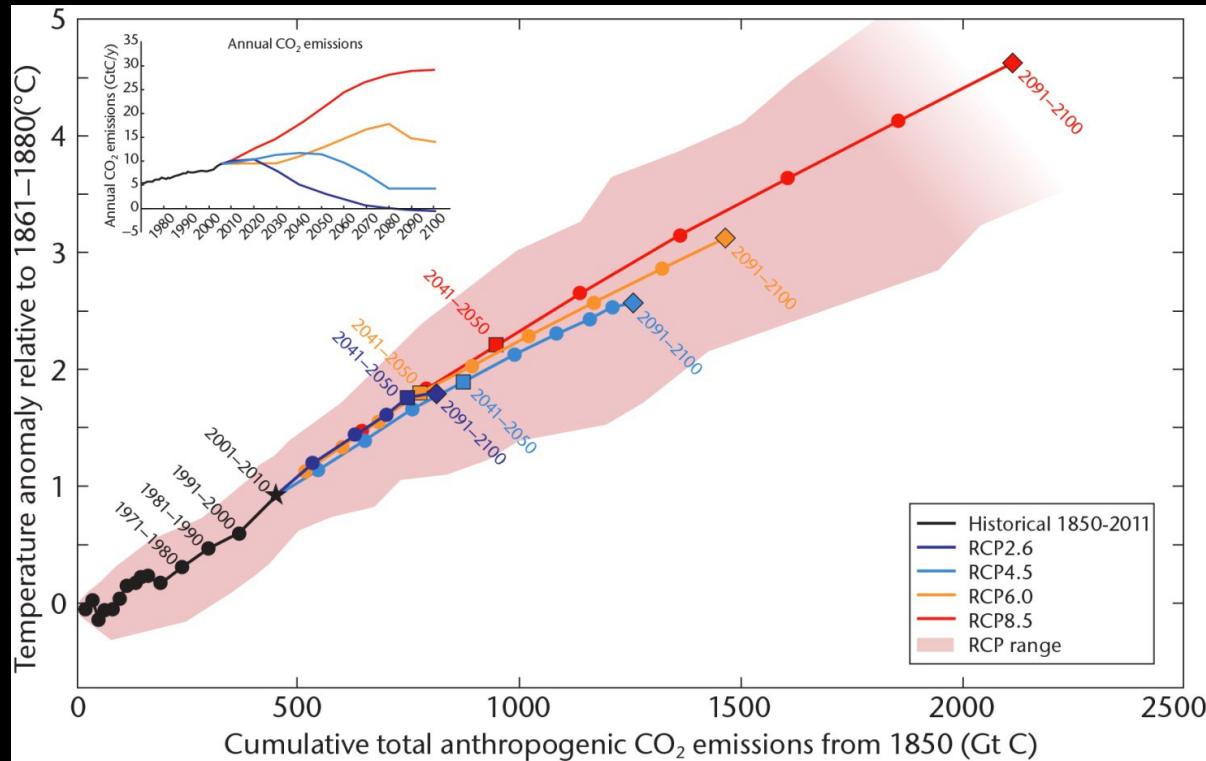
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*UNFCCC, Bonn, 24 October 2013*

# Introduction

- Why model the global carbon cycle?
  - Motivation from climate perspective
- CMIP5 / IPCC AR5 WG1 results
  - “compatible fossil fuel emissions” to achieve a scenario/target
  - Land/ocean model uncertainty
- What's in / not in the models?
  - Processes
  - Ecosystems
  - Model evaluation

# Total CO<sub>2</sub> emissions are strongly linked to total warming



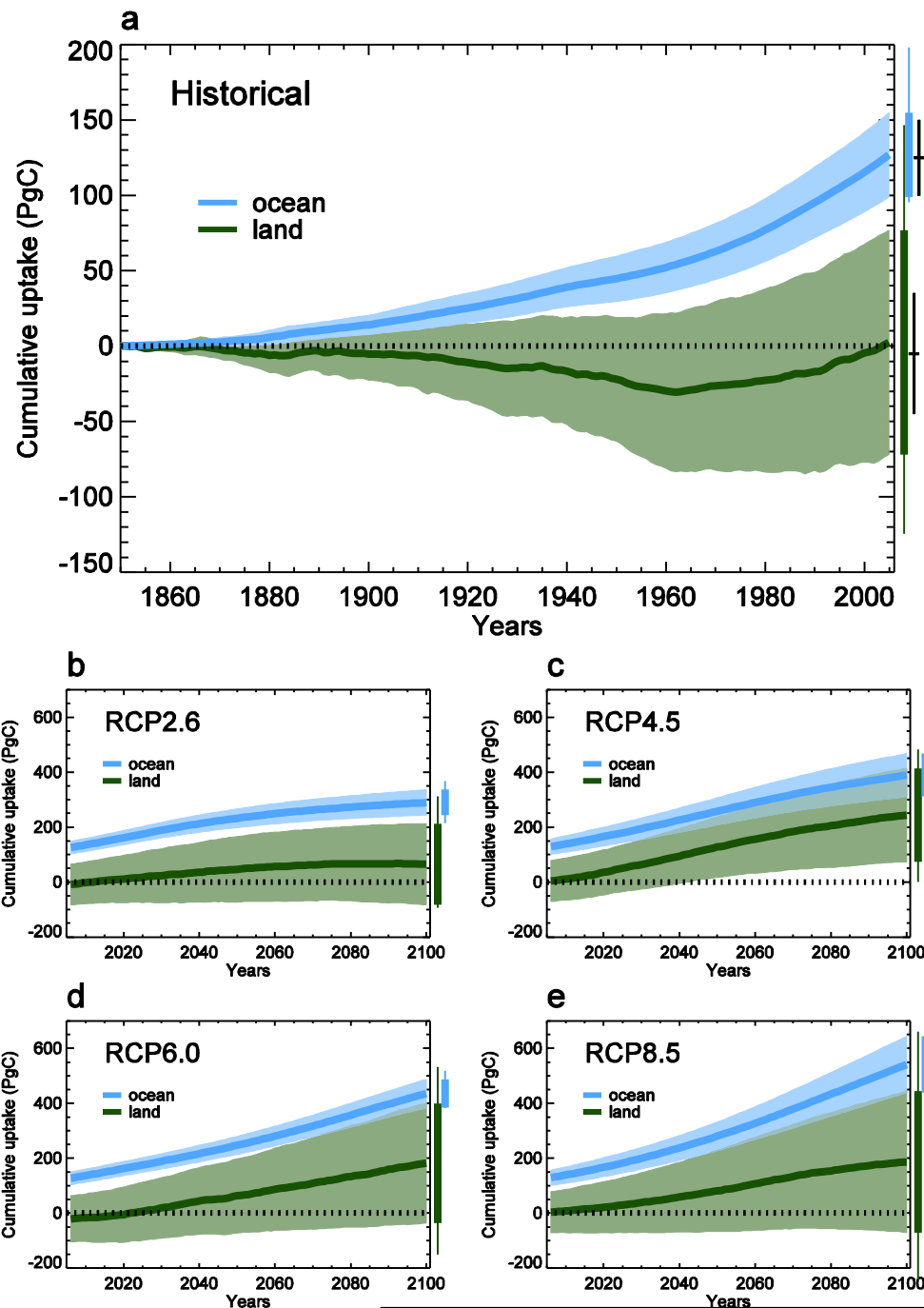
- ⌚ Total CO<sub>2</sub> emissions to 2011 are in range 460-630 GtC.
- ⌚ If warming due to human emissions alone is to be limited to 2°C, total emissions need to be limited to 800-1000 Gt C. More than half of this was already emitted by 2011.

# 21<sup>st</sup> Century global carbon uptake: CMIP5 results

Ocean models agree much better – ocean will continue to take up carbon. For all scenarios. *High confidence.*

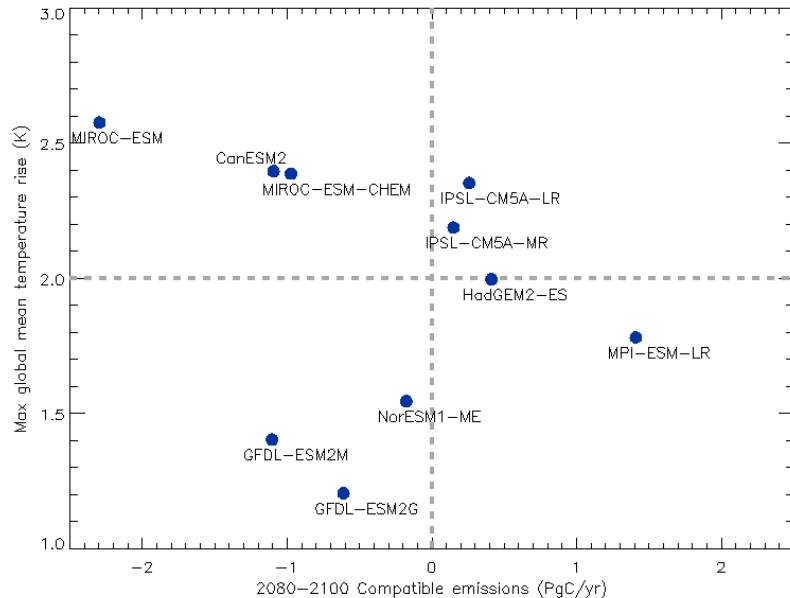
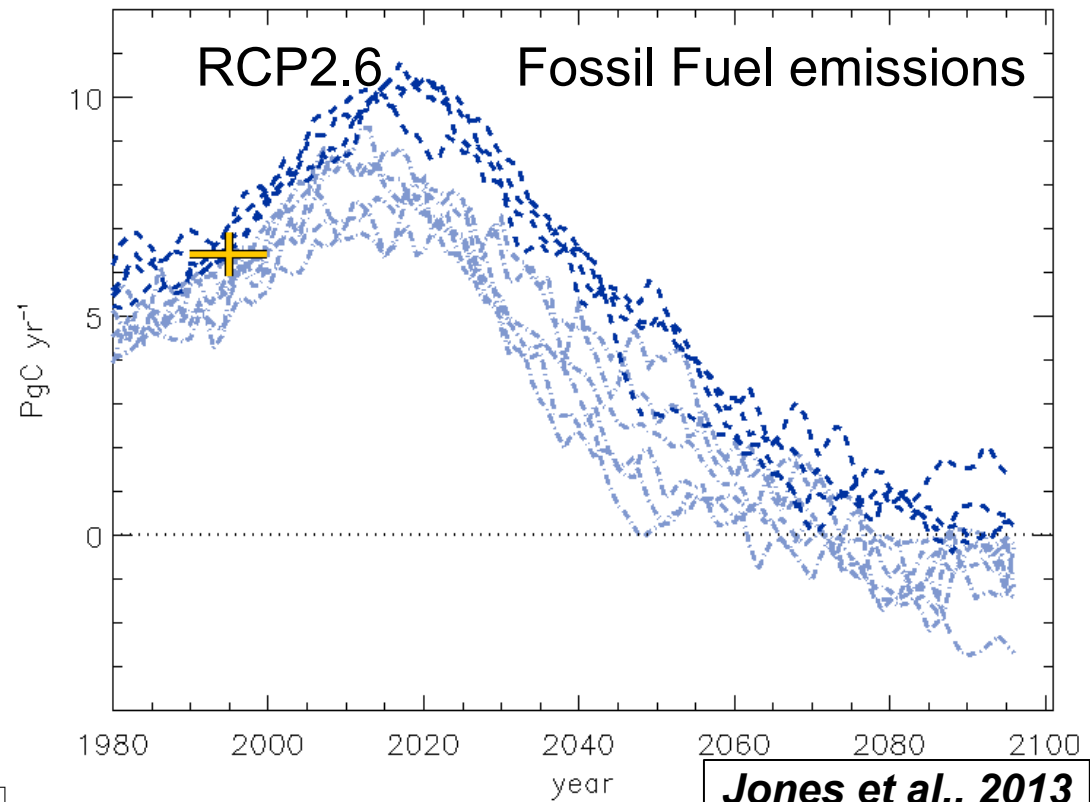
Land models have much greater spread. And also lack some key processes. Therefore *low confidence.*

- But still generally show continued uptake. Minority of models have negative carbon balance, primarily due to land use change.



Simulating the land and ocean carbon uptake allows us to infer CO<sub>2</sub> emissions required to follow scenarios

About ***as likely as not*** that we'll need -ve emissions to achieve RCP2.6

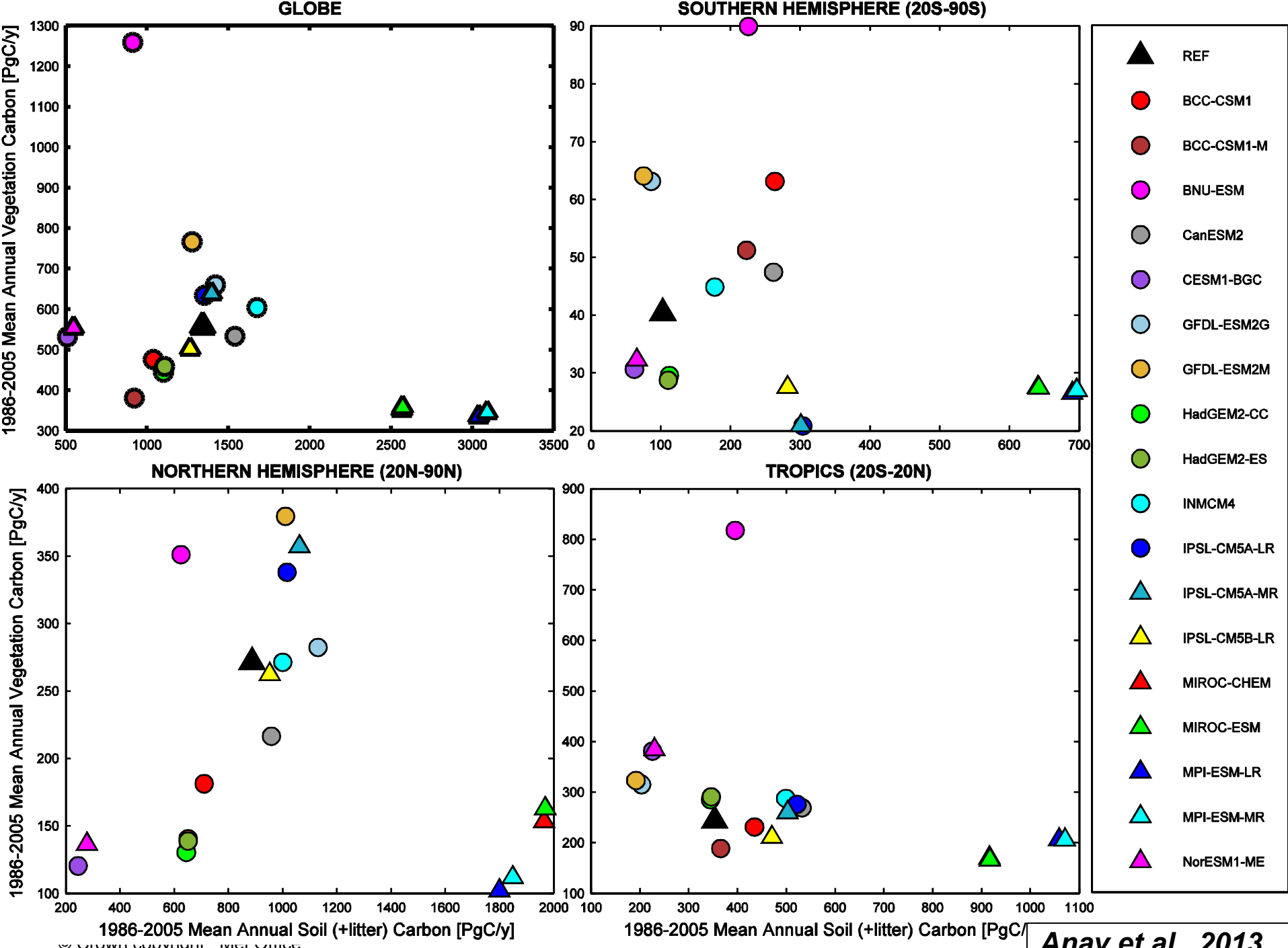


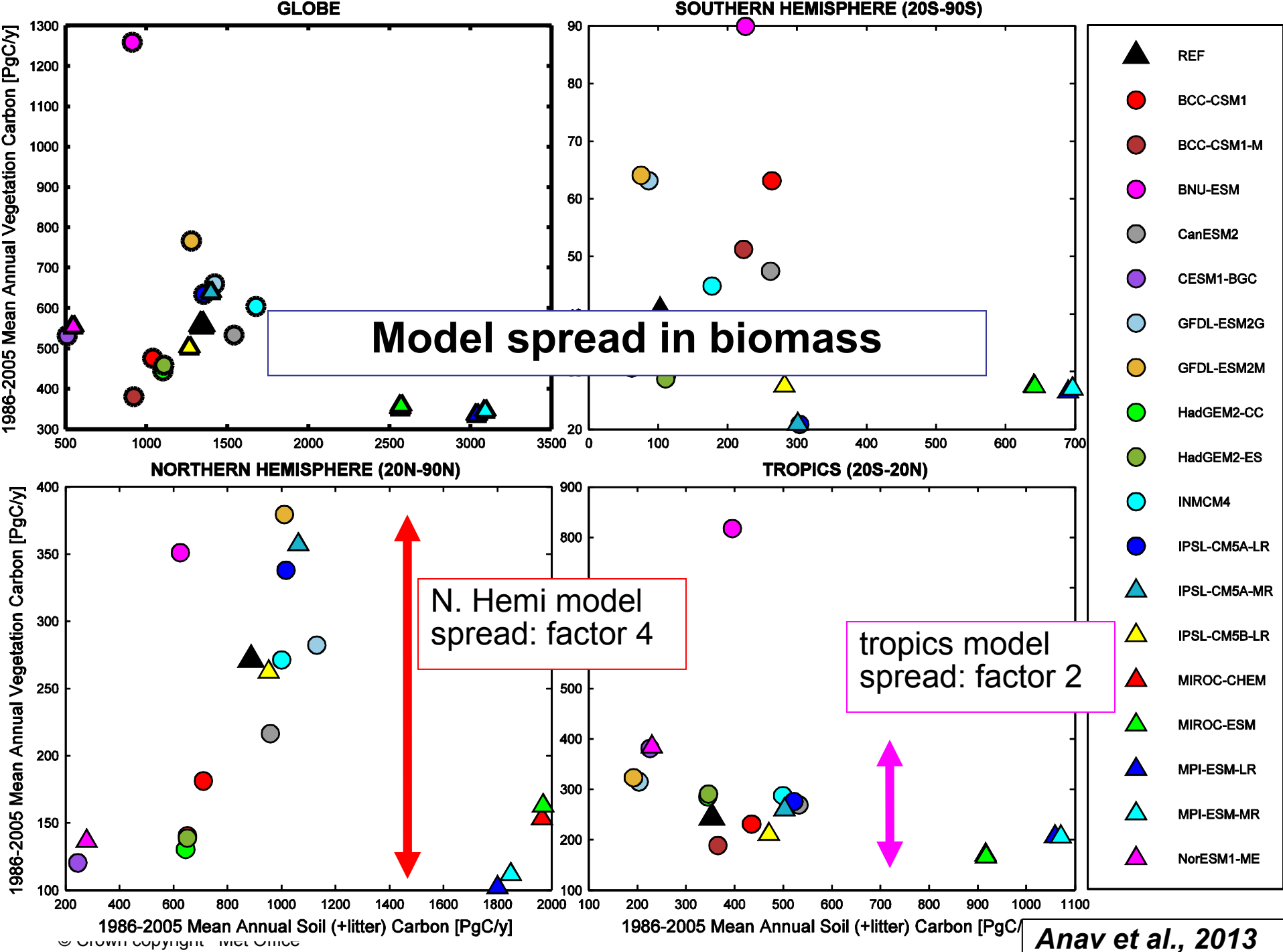
In some models, but not all, this will succeed in avoiding 2 degrees warming

BUT: Any additional carbon source increases the requirement for negative emissions

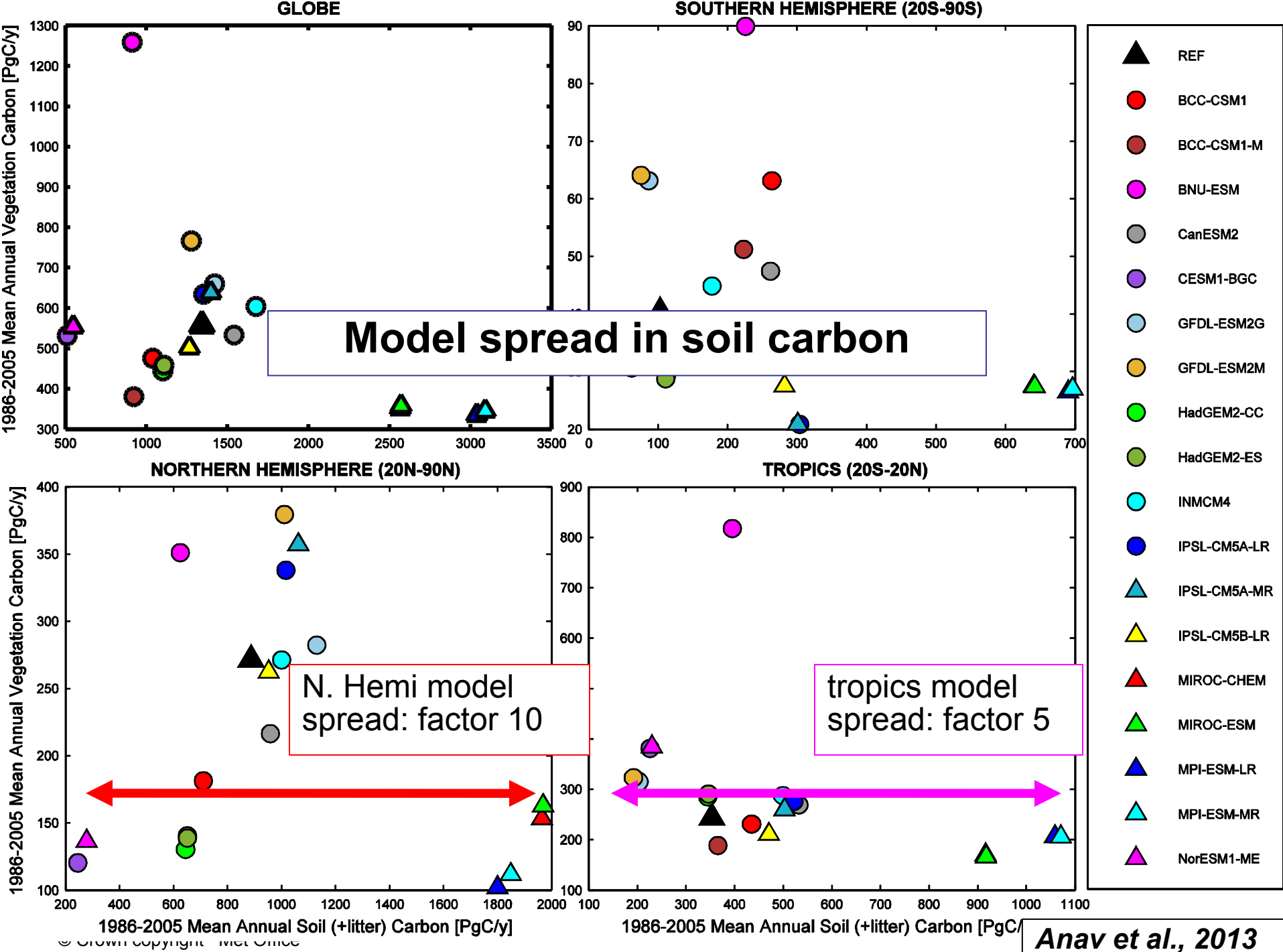
# What's in these models?

- Vegetation functional types
  - Typically 5-10
  - Broadleaf/needleleaf, deciduous/evergreen tree
  - C3/C4 grass
- Soil carbon – SOC in mineral soils (top 1m)
- Need for better model evaluation
  - Big spread in land models
    - Including in initial conditions (global soil and biomass carbon stores)











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# What's missing?

- Missing processes
  - nutrient cycles – nitrogen, phosphorous
- Non-represented ecosystems
  - Permafrost
    - Physics (soil freezing)
    - but not C storage
  - Peatlands/wetlands
    - Vegetation cover – mosses, sedges
    - Deep organic soils
  - Mangroves, seagrass, marshes

**Potential further  
sources/sinks of  
greenhouse gases**

- [not covered further here – but not in global models at all]

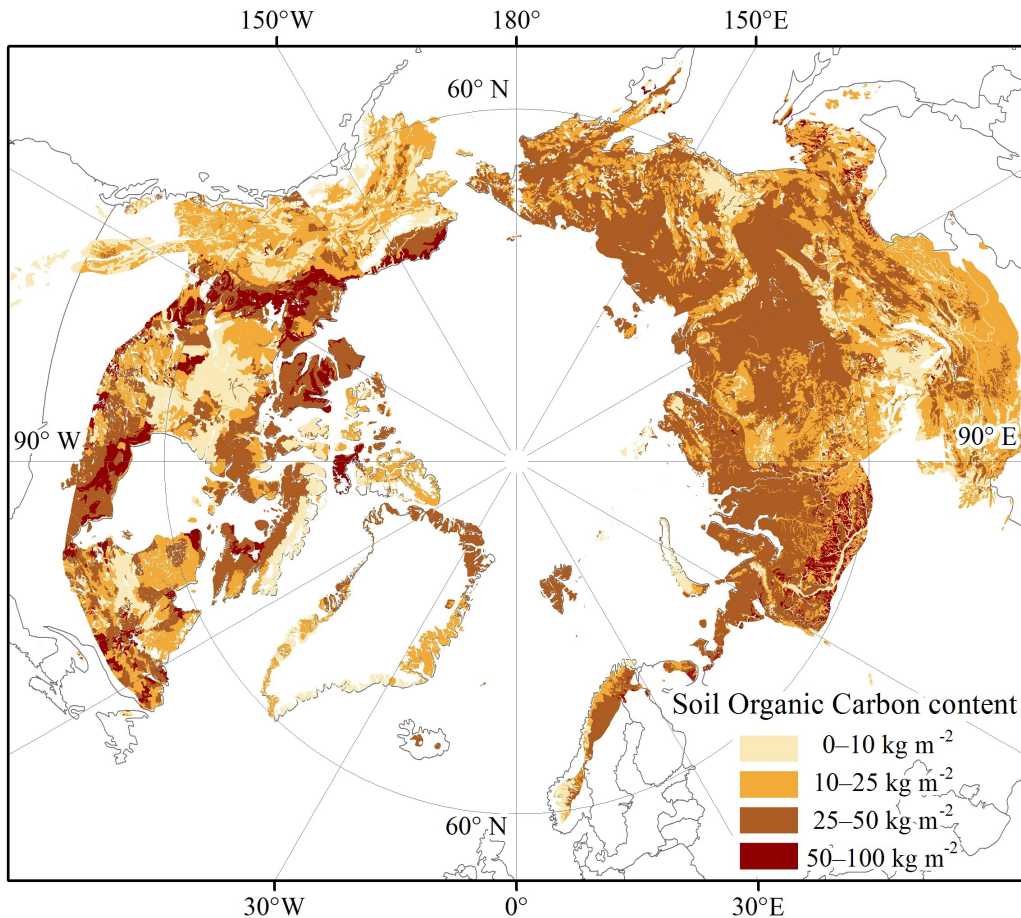
# Observed northern hemisphere permafrost distribution



Permafrost impacts ~ 25 % of the land surface.

Sub-sea permafrost also exists

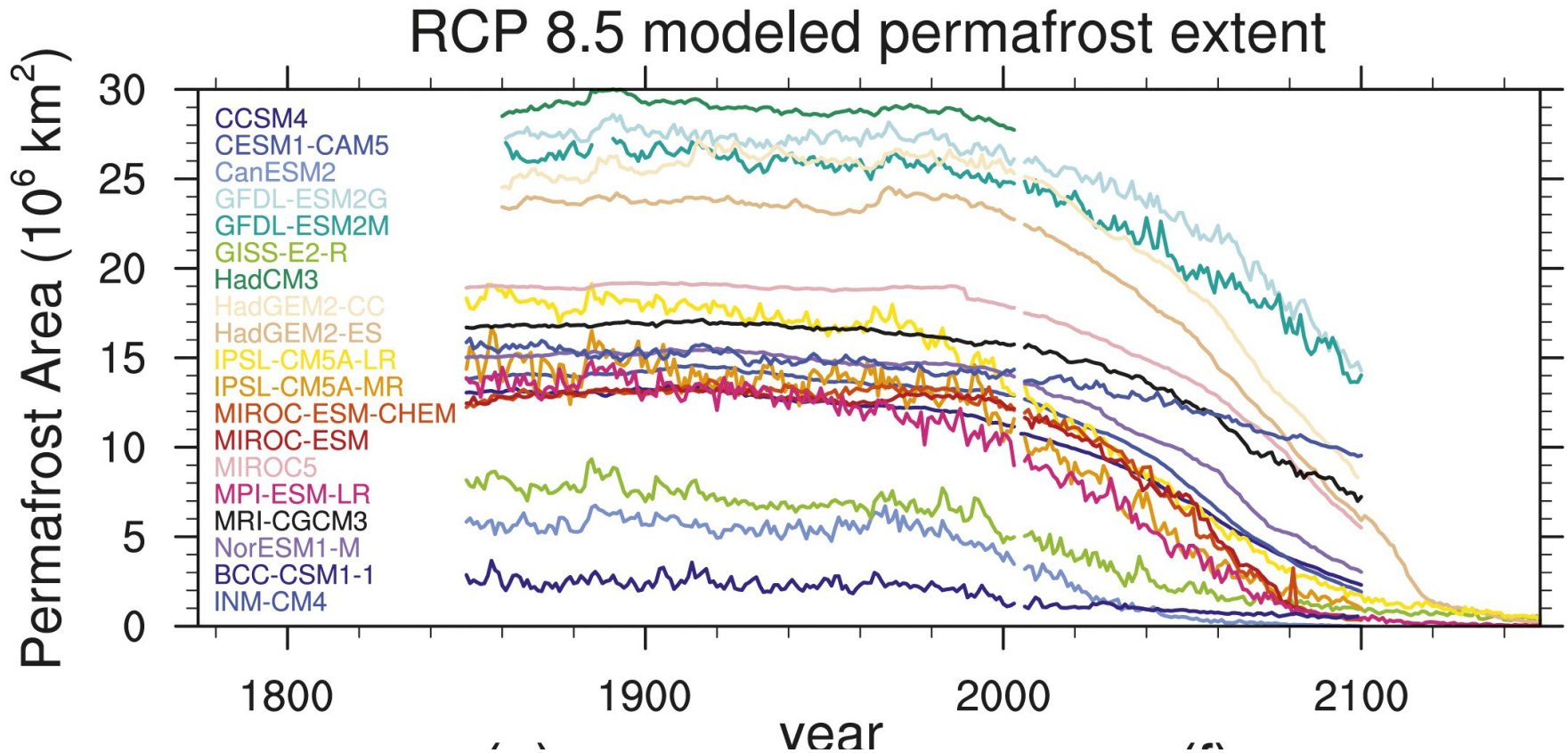
# Frozen soil carbon stores



- Permafrost soils contain around 1,672 Pg carbon (highly uncertain)
- More than twice as much as is currently in the atmosphere.



# Representation of loss of permafrost extent in CMIP5 models



From Koven et al., 2013

“It is **virtually certain** that near-surface permafrost extent at high northern latitudes will be reduced as global mean surface temperature increases.”  
IPCC 2013 [37-81% loss by 2100]

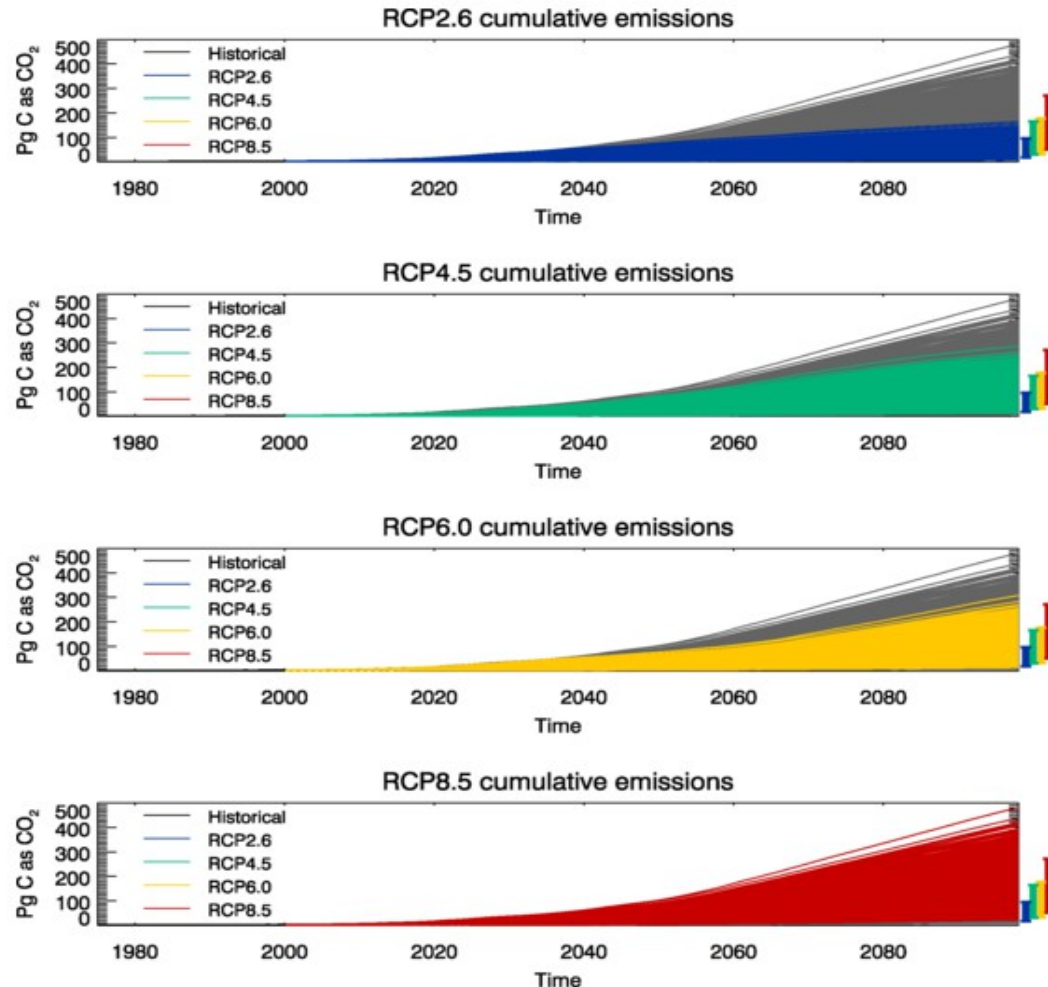


# Soil carbon decomposition

AR5 WG1 assessed a range of 50-250 PgC release, by 2100, for RCP8.5

Uncertainty assessment  
of future carbon release  
from permafrost

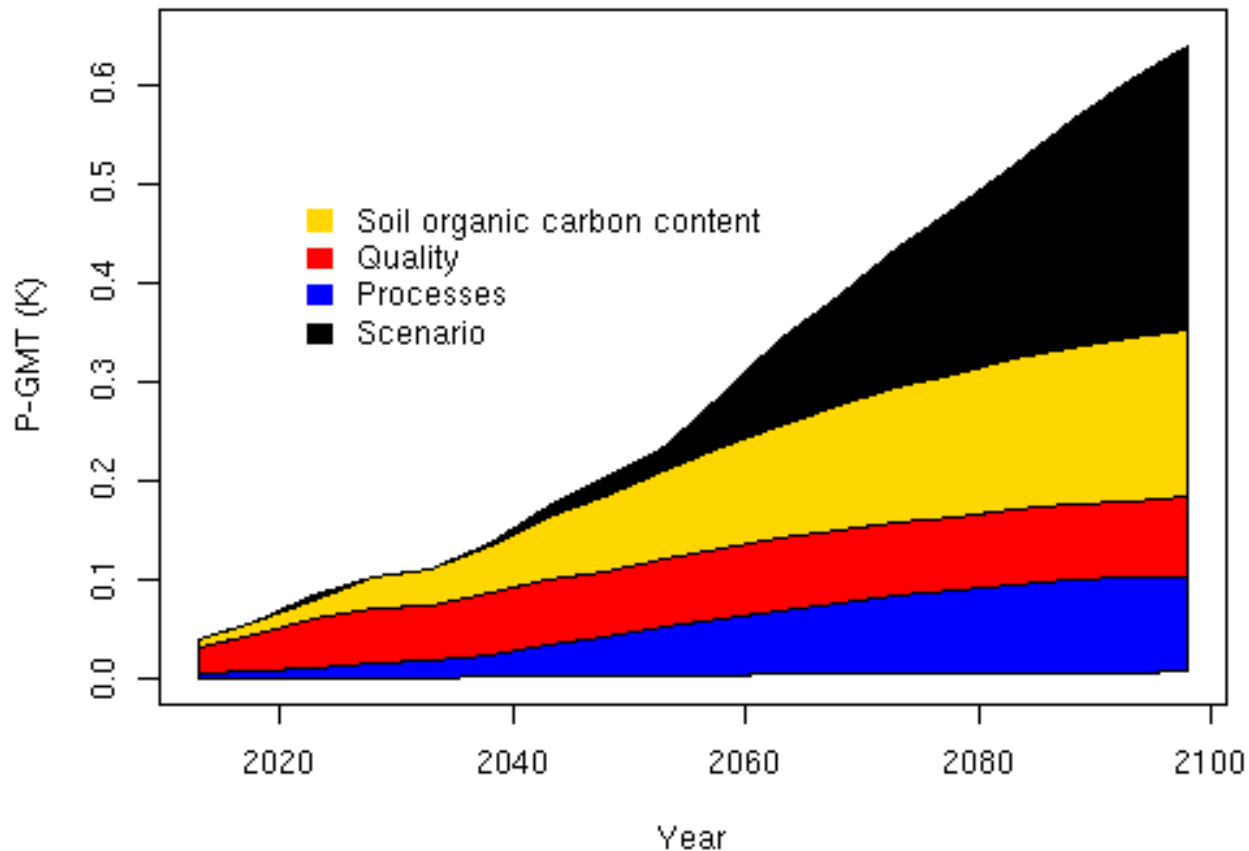
- degree of warming (scenario)
- amount/distribution of carbon
- process uncertainty





# Relative roles of uncertainty

Relative contribution to P-GMT



- PF carbon content and distribution (yellow) one of the biggest sources of uncertainty

- “quality” / lability (red) also important

- Both reducable by better observations



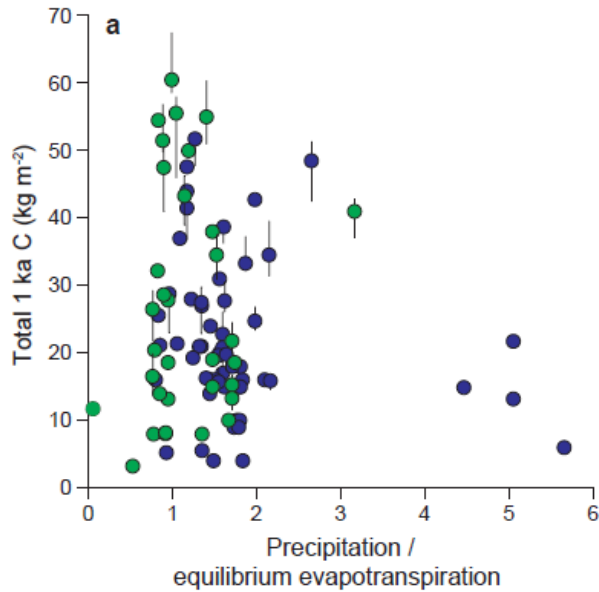
# Peatlands and wetland ecosystems

- The carbon pool in peatlands is big: >600 PgC globally
  - determining controls on future dynamics is an important part of carbon cycle feedbacks to climate change
- Peat accumulation is slow
  - May be dominated by NPP changes rather than decay; spatial and temporal patterns of accumulation over the past millennium (and whole Holocene) suggest a negative feedback – but small.
- Peatland extent may be important and a positive feedback
- Methane emissions clearly important too

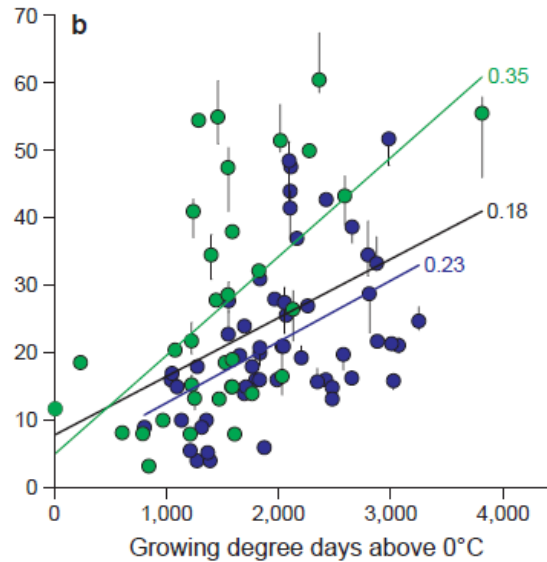


# Spatial variability in total 1kyr peat C

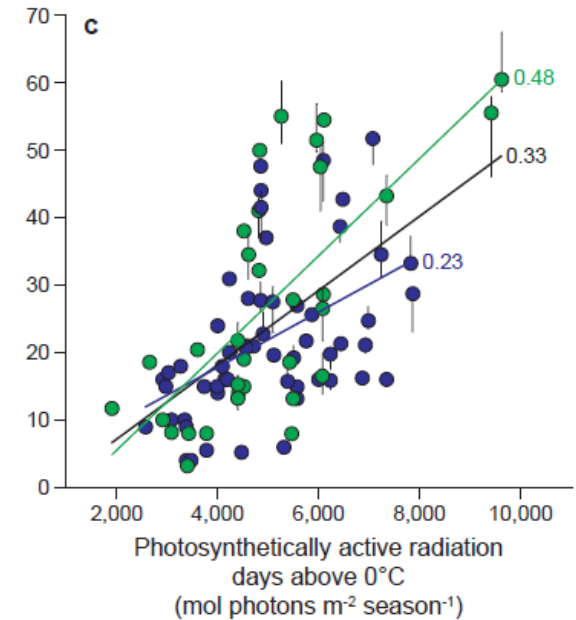
Moisture index



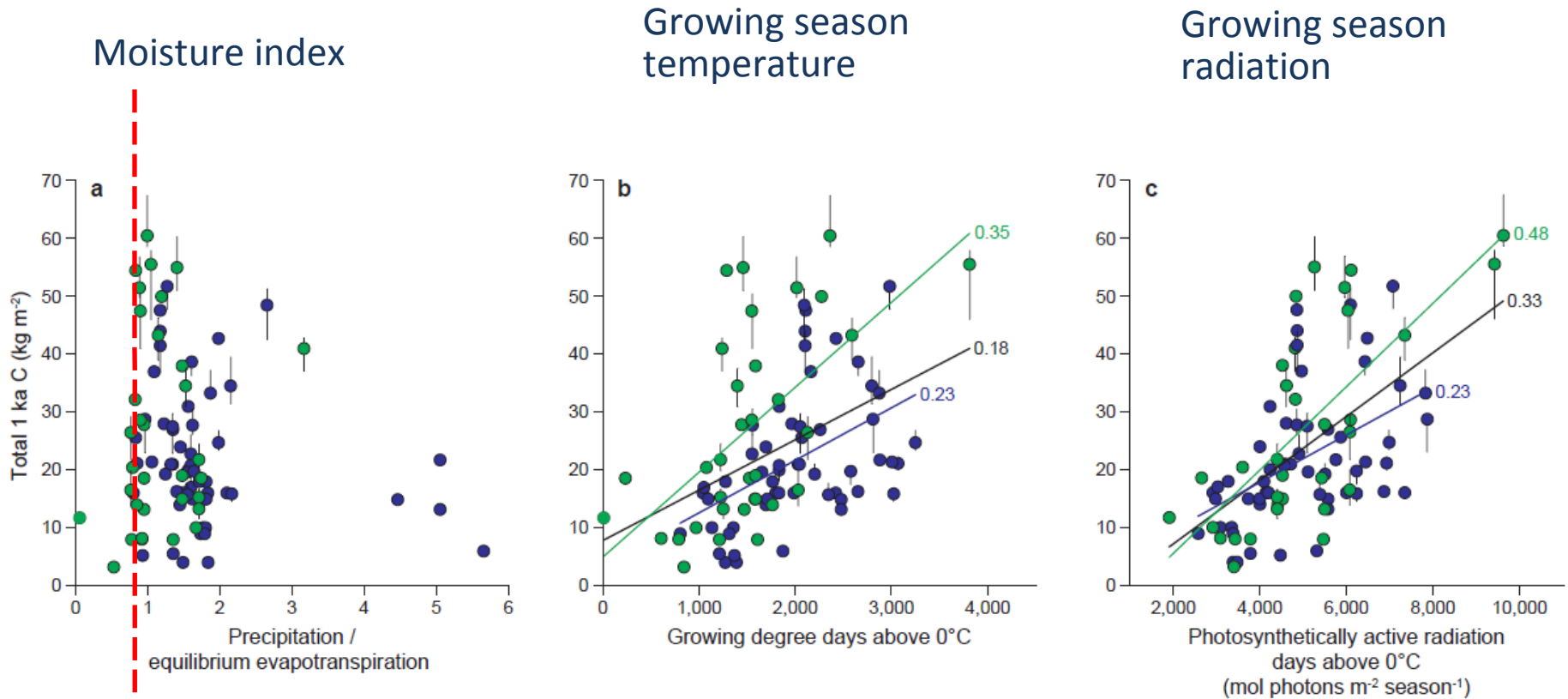
Growing season temperature



Growing season radiation



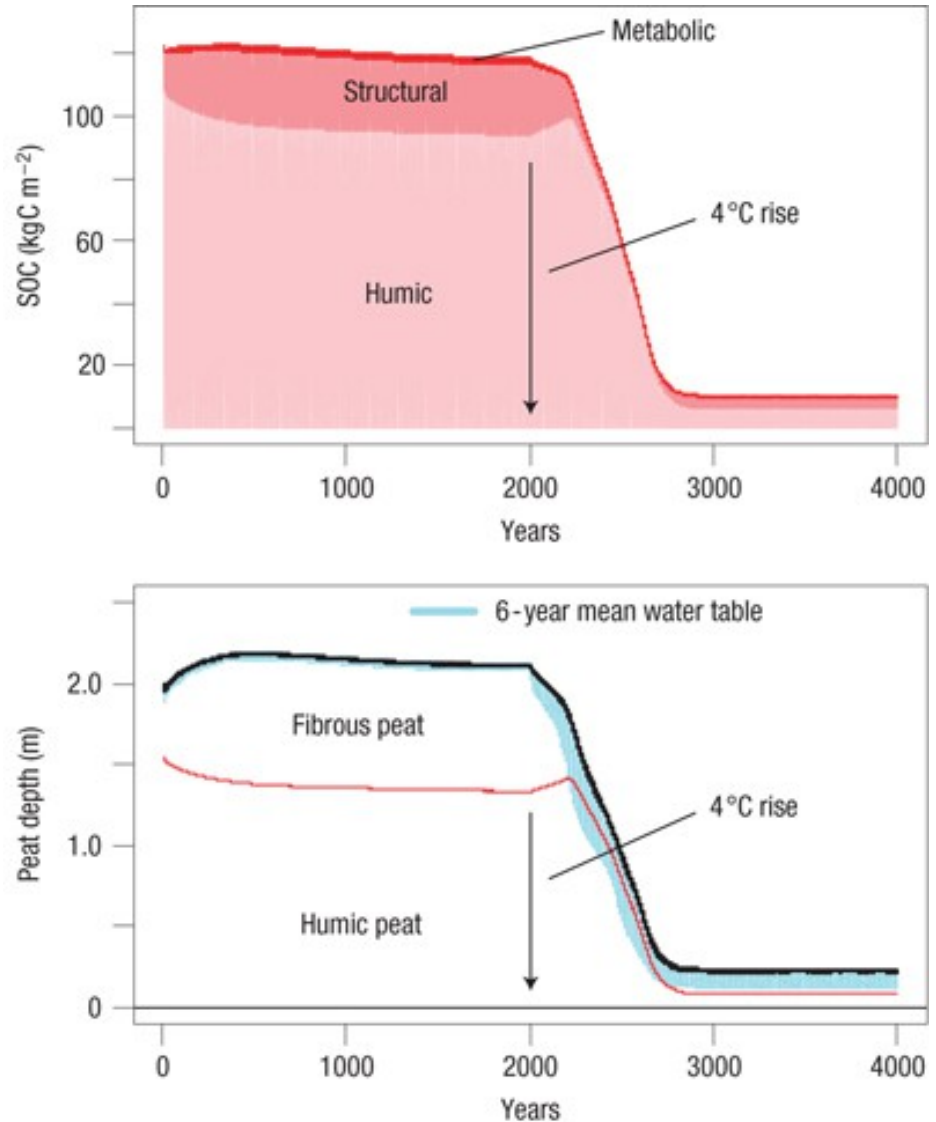
# Spatial variability in total 1kyr peat C



Possible threshold – climate change could drive loss of peatland areas

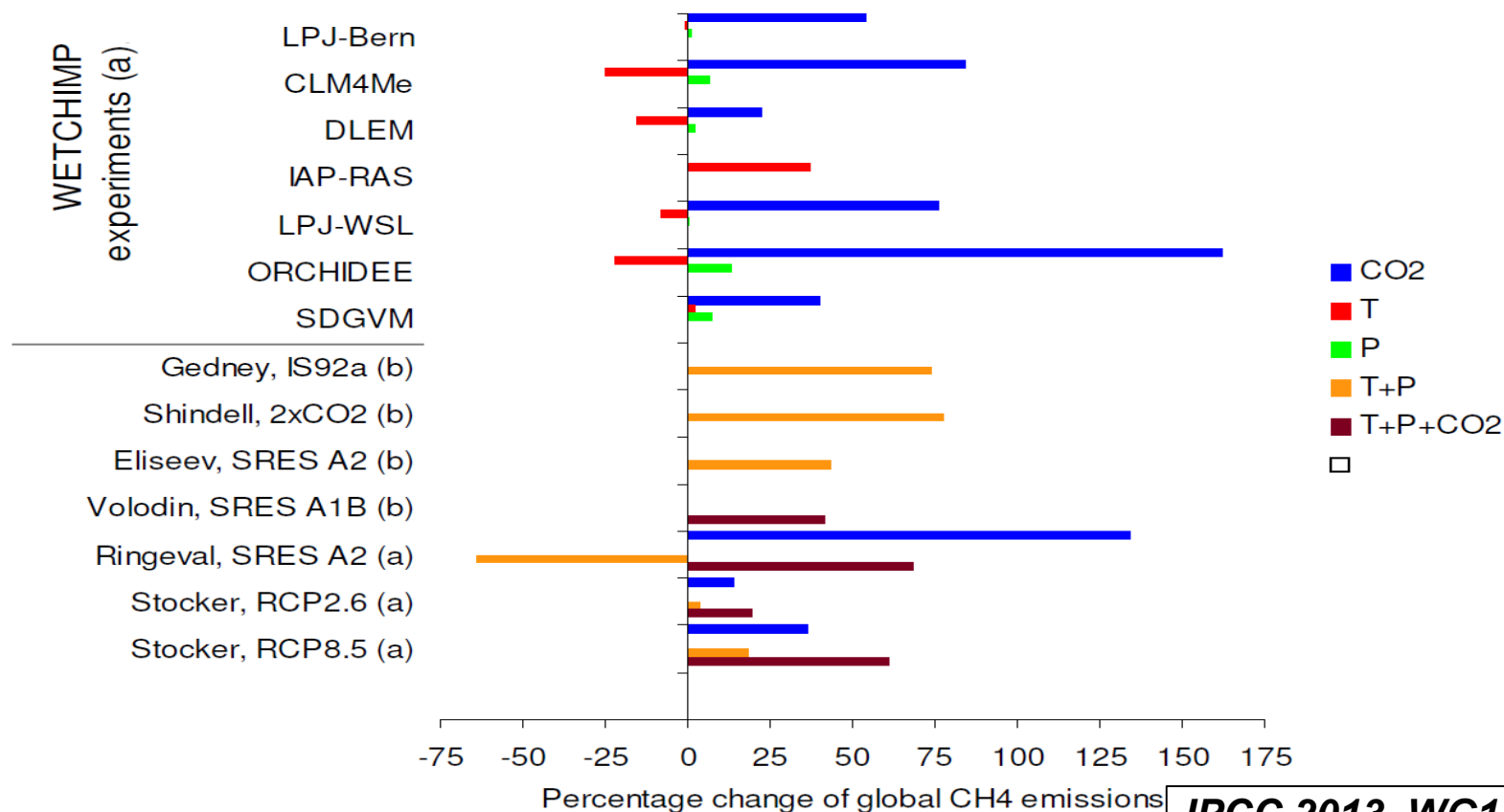
# Positive feedbacks on soil texture may accelerate any peat loss

- Feedback between the water table and peat depth increases the sensitivity of peat decomposition to temperature
  - intensifies the loss of soil organic carbon in a changing climate.
- Peatlands may quickly respond to warming this century by losing labile SOC during dry periods.



# Methane emissions

- Permafrost thaw may also release methane
  - May be more important driver of climate change on short (decade) timescales than CO<sub>2</sub> release
- Wetland methane emissions also **likely** to increase
  - Driven by CO<sub>2</sub> fertilisation as well as a changing climate





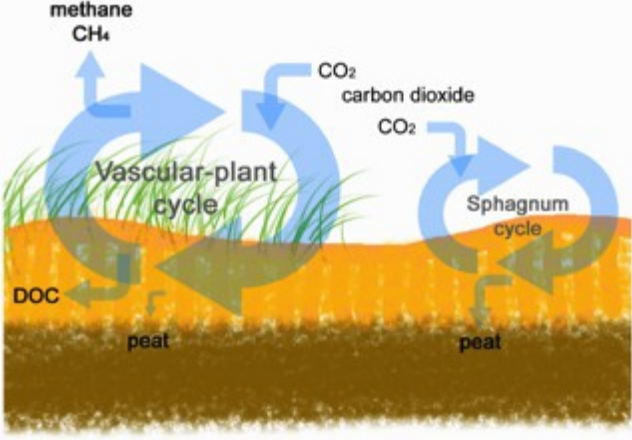
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# Summary

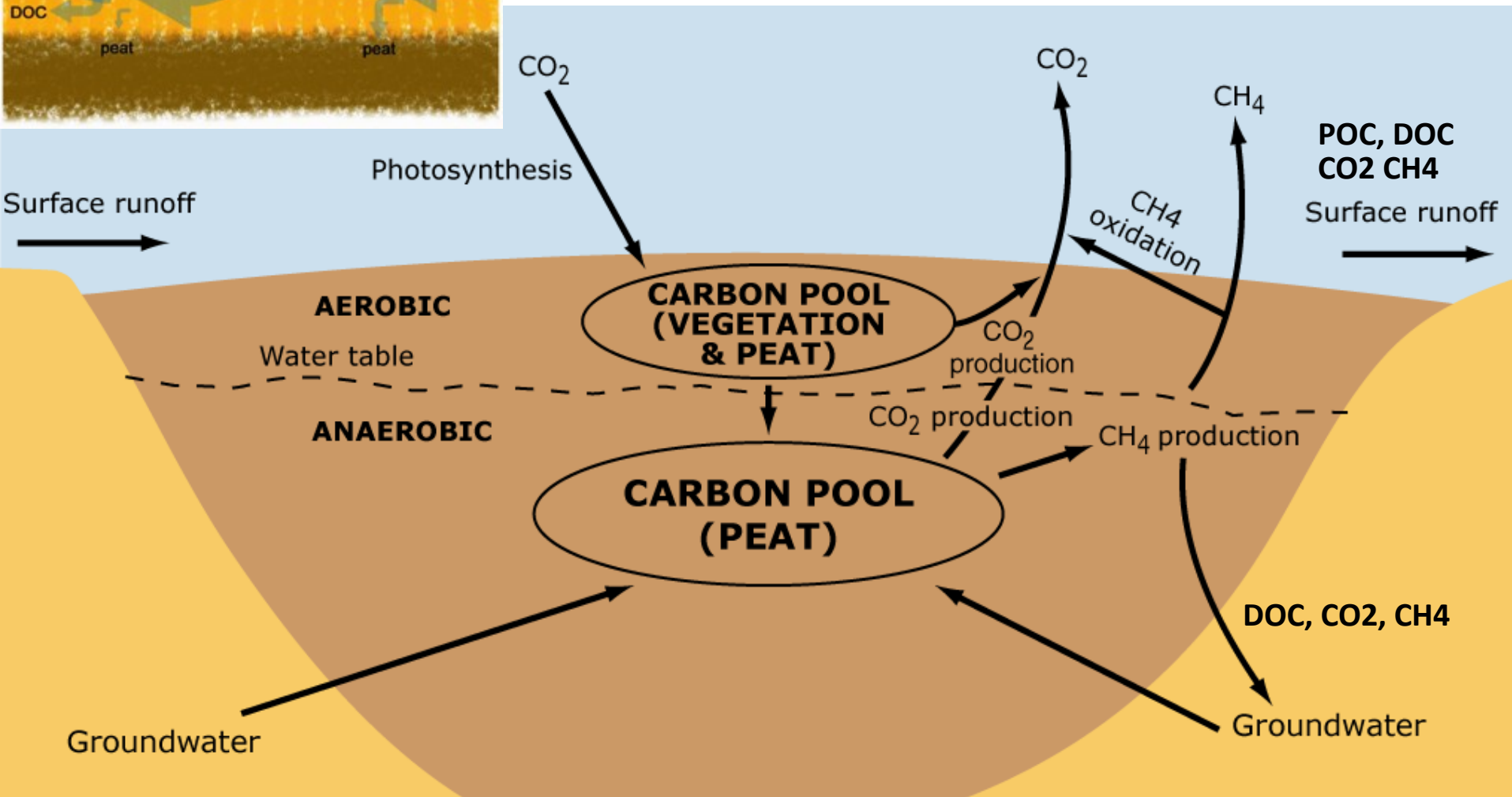
- Global carbon cycle responds to climate change
  - Essential link between human activity and climate change
- Low mitigation CO<sub>2</sub> pathways, such as RCP2.6 as likely as not require global negative emissions
  - But large model uncertainty (especially on land)
  - Missing process/ecosystems play big role in achievability of targets
- Need better / more complete inventories:
  - Model evaluation / initial conditions
  - Ecosystems
    - Permafrost
      - Likely to release additional carbon this century
    - Peatlands/wetlands
      - May release additional carbon (CO<sub>2</sub>) this century
      - Likely to emit more methane this century
  - Both provide additional positive feedback onto climate change

# References

- Anav et al., 2013, J. Climate (*C4MIP Special Collection*)
- Burke et al., 2013, The Cryosphere
- Ise et al., 2008, Nature Geoscience
- Jones et al., 2013, J. Climate (*C4MIP Special Collection*)
- Koven et al., 2013, J. Climate (*C4MIP Special Collection*)



# Peatland carbon cycling



Higher water table: increased  $\text{CH}_4$  production, reduced respiration